

Computer Programs for Avalanche Runout Prediction

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COMPUTER PROGRAMS FOR AVALANCHE RUNOUT PREDICTION

by

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ABSTRACT

This report summarizes the function and application of three computer programs, made operational on the Shinjo Branch computer system, for use in analysis of snow avalanche runout. In the use of fluid mechanics principles to model snow flow, two of the codes are based upon equilibrium hydrodynamic equations, while a third incorporates transient, viscous effects in a two-dimensional incompressible boundary layer formulation. Of the two hydrodynamic based codes, one which was previously developed, has constant frictional and viscous material coefficients that vary significantly with different avalanche types, which makes it difficult to apply, except by experienced persons. The second code, developed at Shinjo Branch, brings in a flow depth dependence and a material locking property in definition of the material coefficients, which reduces significantly the range on the coefficient values for different avalanche types. Listing of each code is included in this reporting, as is the format and order in input data preparation. Comparison of predicted velocity profiles and runout distances from each code is made for one avalanche path (Ironton Park, Colorado USA). While velocity profiles are different for each code, runout distance can be matched by selective choice of parameters.

Introduction

In this report is summarized the results of an investigation having the following dual purposes:

1. To make operational on the Shinjo Branch computer, certain codes that are used in analysis of snow avalanche runoff.
2. To develop and checkout a modified version of a computer code, that is based upon equilibrium flow dynamics, but incorporates recent developments in the mechanics of avalanche flow.

The codes that have been made operational on the Shinjo Branch computer (Melcom 70 Computer System) are:

1. Program AVALNCH — This program models the two-dimensional transient flow of a viscous fluid. The code has been used to analyze numerous avalanche paths and different avalanche types. The code was developed by imposing restrictive conditions, unique to avalanche dynamics, upon a general purpose fluids code. The general purpose code can be used for a wide range of transient viscous fluid problems, including impact dynamics. Two versions of program AVALNCH are considered in this reporting. One version, operational since 1978, uses a so called "fast-stop" option to model the slow-down of avalanches at low-speed terminal flow. This modeling is necessary because of thixotropic character of flowing snow, which has a tendency to lockup as the flow speed reduces to a stop condition. The fast-stop algorithm is an empirical representation of the locking property. A modified version of AVALNCH, which incorporates a biviscous modeling of snow, is reported also. The program was developed during the course of this reported work. The biviscous representation of snow, approaching the Bingham fluid idealization of a locking material, is a more physically based approximation

of the snow locking property than fast-stop. Both versions of program AVALNCH use two parameters to represent the fluid state of flowing snow.

2. Program ACCEL — This program, developed by Cheng and Perla (1979), is based upon an equilibrium viscous fluid modeling of avalanching snow. The program uses hydrodynamic equations of two-dimensional flow, in which two parameters relating to surface friction and viscous drag are selected for an avalanche analysis.
3. Program BIEQ — This program, developed during the course of this work, is based upon the fluid dynamics of viscous equilibrium flow. It is a modification of program ACCEL in which recent mechanics principles of avalanching snow are incorporated, at least to some degree of approximation. The program incorporates a two-parameter material representation, but also accounts for material locking, explicitly.

In the following reporting of these computer codes, some details on the use of the codes are given, as well as results of analysis of avalanching snow. Program results are compared, and conclusions drawn from these comparisons.

COMPUTER PROGRAM AVALNCH

Computer program AVALNCH is a specialized version of a general purpose program called SOLA-SURF, which was developed by Hirt, Nichols and Romero (1975) at Los Alamos Scientific Laboratories. Program SOLA-SURF models the 2 dimensional transient flow of an incompressible viscous fluid which may have a free surface. The programs are based upon numerical integration, using finite difference methodology, of the 2 dimensional equations of motion of a viscous fluid. The equations are

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} = g_x - \frac{1}{\rho} \frac{\partial p}{\partial x} + \nu \nabla^2 u$$

$$\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} = g_y - \frac{1}{\rho} \frac{\partial p}{\partial y} + \nu \nabla^2 v$$

where $\nabla^2 = \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial y^2}$

In addition to these equations, the equation of conservation of mass, namely

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} = 0$$

also enters the computations. Computations are carried out in two steps within the computer code. At each cycle (CYCLE) the fluid is advanced in the grid, based upon the gravitational driving force and the frictional drag force acting on each cell that contains fluid. Following this calculation, the fluid is redistributed by one or more iterations (ITER) in order that the total mass of material not change within a specified limit of accuracy. At the start of each cycle Eqs (1) are solved, whereas in the iterative phase simple linear equations are used, since the change needed in order to conserve mass, is small.

Program AVALNCH has been specialized to model the flow of a snow avalanche over an extended path by several simplifying assumptions that significantly reduce computer running time, compared to that of a general fluid modeling of the problem. The basic simplifications are:

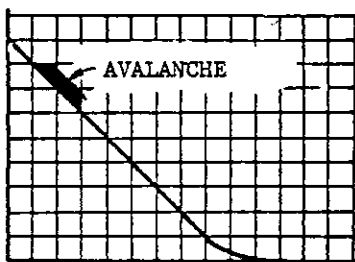
1. Specialization of the avalanche flow to that of boundary layer

flow, for which a single vertical cell is used to represent the depth of the flow.

2. Representation of the actual avalanche path profile by a horizontal grid of elements, for which slope-parallel and slope-normal gravity components are specified.

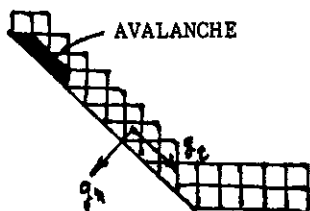
These simplifications can be demonstrated graphically by a sequence of diagrams (Figure 1). It is representation (c) of Figure 1 that is used in program AVALNCH. This approximation excludes only impulses impressed upon the flowing material as the profile slope decreases. However, the effect of this has been shown to be negligible for ordinary avalanche paths (Cheng, Perla, 1979). The reduction to two cells in the vertical direction, has resulted in significant reduction in computer running time; yet the results show accurate prediction of avalanche speeds and runout distances along the path. In using only two vertical cells, the representation of any vertical variation in the flow parameters is excluded, so that modeling of flow depth should not be expected to be accurate. To represent vertical effects more accurately, additional vertical cells should be used (and may be used in program AVALNCH), but, with the long runout distances of avalanches, the computer cost may become excessive.

A number of avalanches have been modeled with program AVALNCH, so that the range in the basic parameters of the code have been established (Lang, et al., 1979), (Martinelli, et al., 1980). The two parameters are viscosity, ν , and friction coefficient, f . In the case of high altitude mid-winter, strongly sintered, dry snow avalanches, as occur in the Rockies, the values of f and ν that model the flow are numerically equal at $f=0.45$ and $\nu=0.45\text{m}^2\text{s}^{-1}$. For low altitude coastal wet snow avalanches, values as high as 0.6 to 0.8 have been used. For weakly sintered dry snow avalanches values as low as 0.4 have yielded adequately modeled results.



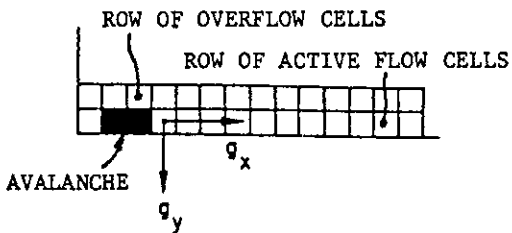
(a)

HERE THE AVALANCHE PROFILE IS SUPERIMPOSED UPON A RECTANGULAR GRID OF CELLS - THE SETUP FOR THE ORDINARY APPROACH TO NUMERICAL FLOW PROBLEMS



(b)

HERE IF SHOWN A GRID SETUP FOR A BOUNDARY LAYER FLOW PROBLEM IN WHICH ONE GRID DIMENSION IS LONG - BUT WITH THIS SETUP BOUNDARY CONDITIONS ARE COMPLEX



(c)

HERE IS SHOWN A HORIZONTAL GRID REPRESENTATION OF THE FLOW PROBLEM. GRAVITY COMPONENTS g_x AND g_y ACCOUNT FOR THE DRIVING AND CONTACT FORCES OF THE MOTION, WHICH MAY VARY WITH EACH CELL

FIGURE 1: FLOW DOMAIN REPRESENTATIONS

It is suggested that if the program is to be used in site specific applications that test cases be run to determine the range of values of f and v .

AVALANCHE ANALYSIS USING PROGRAM AVALNCH

Several steps are involved in setting up a problem to be run with program AVALNCH. Once an avalanche path is selected, the first step is to draw a profile of the path using the same scale in the vertical and horizontal directions. Generally, data from which a profile is drawn is taken from topographic maps of the avalanche region. A typical profile plot of an avalanche path in Colorado (Ironton Park avalanche path) plotted from a 1:25,000 topographic map is shown in Figure 2. The profile may be approximated by a continuous curve as in Figure 2, or by a series of straight line segments, which is computationally easier. Having drawn the profile, the next step is to lay off a uniform grid along the slope, selecting the grid dimension so that less than 200 grid lines are used along the path. In the case of Ironton Park 110 grid lines were used, separated by 10.0m increments. By some measurement or calculation technique the change in elevation from one end of each 10.0m element (or cell) to the other end must be determined. For example, for Ironton Park the elevation change of each element was determined by rule measurements on the profile (Table 1). Accuracy of measurement of each element elevation change is not as important as having the total elevation change equal that of the profile, and this should be checked each time a profile is set up. For example, in the case of Ironton Park the elevation changes in Table 1 could be rounded off to whole numbers and not significantly change the velocity along the path, provided the total elevation change remains the same.

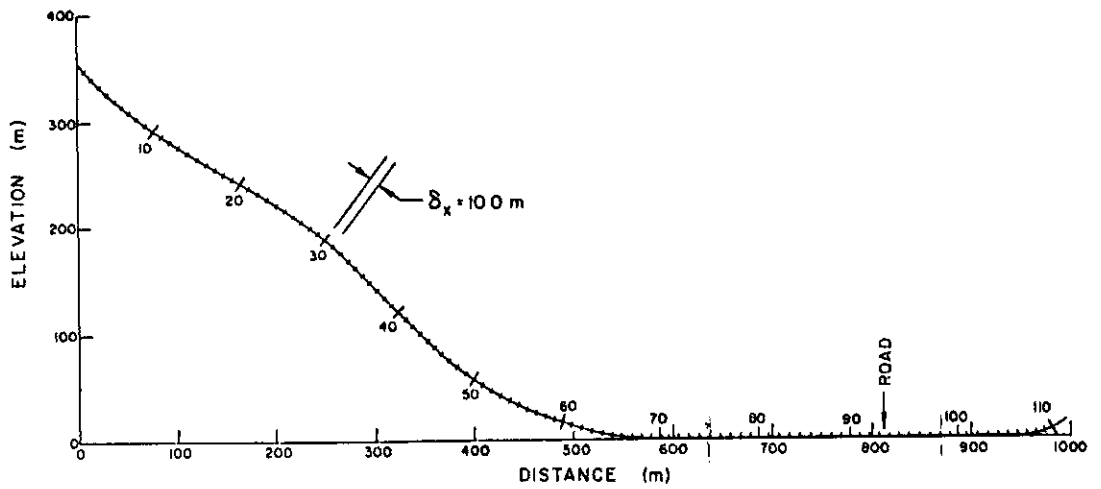


FIGURE 2: IRONTON PARK AVALANCHE PATH PROFILE

Table 1: Ironton Park element elevation change

ELEMENT	ELEVATION CHANGE (m)								
1 to 8	6.5	6.2	6.2	6.2	6.5	6.2	6.2	6.2	6.2
9 to 16	6.2	6.5	6.5	6.2	5.8	6.0	5.8	5.6	
17 to 24	5.6	5.2	4.6	4.2	4.2	4.6	5.2	4.6	
25 to 32	4.8	5.0	5.2	5.4	5.6	5.8	6.2	6.2	
33 to 40	6.2	6.2	6.5	6.8	7.4	7.2	7.2	7.1	
41 to 48	7.0	7.0	7.0	6.6	6.2	6.2	6.0	6.0	
49 to 56	5.6	5.2	5.2	5.0	4.6	4.2	4.2	4.0	
57 to 64	4.0	3.6	3.6	3.4	3.4	3.0	2.8	2.5	
65 to 72	2.2	1.8	1.6	1.4	1.2	2.0	1.3	0.8	
73 to 80	0.4	0.1	0.0	0.0	0.0	0.0	0.0	0.0	
81 to 88	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	
89 to 96	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	
97 to 104	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	
105 to 110	0.0	0.0	0.0	0.0	0.0	0.0			

In the Ironton Park example there are no elements exhibiting what is termed adverse slope, which is slope that the avalanche must climb. In specifying adverse slope the elevation change (as in Table 1) is listed with negative values. Having determined the elevation change for each element of the profile, the remaining steps involve preparation of the computer input data.

AVALNCH INPUT FORMAT

Input in the line sequence given below may be submitted either as a data file, or as a sequence of cards depending upon user preference.

Line 1: FORMAT (40A2)

Columns 1-80: Title and identification information

Line 2: FORMAT (2I10, 5F10.0, I10)

- Columns 1-10: IBAR - number of cells in the slope-parallel direction; maximum is 200, unless program is changed.
- Columns 11-20: JBAR - number of cells normal to the path; maximum is 2, unless program is changed.
- Columns 21-30: DX - dimension of cell along path (m).
- Columns 31-40: DY - dimension of cell normal to path (m).
- Columns 41-50: YU - kinematic viscosity (m^2s^{-1}).
- Columns 51-60: FK - friction coefficient; if given zero value here, then must input an array of friction coefficients for each cell (Line 5).
- Columns 61-70: TF - avalanche flow time (s).
- Columns 71-80: NP - number of cycles between extended printouts.

Line 3: FORMAT (8F10.0)

- Columns 1-10: thickness of avalanche slab in cell #1 (m).
- Columns 11-20: thickness of avalanche slab in cell #2 (m).
- Columns 21-30: thickness of avalanche slab in cell #3 (m).
- Columns 31-40: thickness of avalanche slab in cell #4 (m).
- Columns 41-50: thickness of avalanche slab in cell #5 (m).
- Columns 51-60: thickness of avalanche slab in cell #6 (m).
- Columns 61-70: thickness of avalanche slab in cell #7 (m).
- Columns 71-80: thickness of avalanche slab in cell #8 (m).

Must continue this listing on succeeding lines until IBAR entries are specified, including zero-thickness cells.

Line 4: FORMAT (8F10.0)

- Columns 1-10: change in elevation of cell #1 (m).
- Columns 11-20: change in elevation of cell #2 (m).
- Columns 21-30: change in elevation of cell #3 (m).
- Columns 31-40: change in elevation of cell #4 (m).
- Columns 41-50: change in elevation of cell #5 (m).
- Columns 51-60: change in elevation of cell #6 (m).
- Columns 61-70: change in elevation of cell #7 (m).
- Columns 71-80: change in elevation of cell #8 (m).

Must continue this listing on succeeding lines until IBAR entries are specified.

Line 5: FORMAT (8F10.0)

(This set of data is required if FK=0 on Line 2)
Columns 1-10: friction coefficient for cell #1
Columns 11-20: friction coefficient for cell #2
Columns 21-30: friction coefficient for cell #3
Columns 31-40: friction coefficient for cell #4
Columns 41-50: friction coefficient for cell #5
Columns 51-60: friction coefficient for cell #6
Columns 61-70: friction coefficient for cell #7
Columns 71-80: friction coefficient for cell #8

Must continue this listing on succeeding lines
until IBAR entries are specified.

This completes specification of input data for program AVALNCH.

Examples of input data, and of program output for the Ironton Park
avalanche path are published (Lang, etal. 1979).

PROGRAM AVALNCH INTERNAL LOGIC

The original developers of program SOLA-SURF (Hirt etal., 1975) have a complete discussion on the logic of the program, which will not be repeated herein. In summary, the logic is shown by a flow chart of the program (Figure 3). Distinction between a cycle (CYCLE) and an iteration (ITER) is indicated on the flow chart. Section 1000, which pertains to a CYCLE, contains the complete Navier-Stokes equations. Section 3000, which pertains to ITER, contains simplified linear equations for small perturbation of parameters in order to achieve conservation of mass. The boundary condition section 2000, is basically the only section that must be changed in order to apply the program to different problem types. For example, for impact problems, the velocities in cells that represent a barrier must be zeroed in section 2000. However, if forces on a barrier are to be computed, then equations for this should be placed in section 4280 just prior to outputting data for the current CYCLE. In the case of impact problems it is also necessary to work

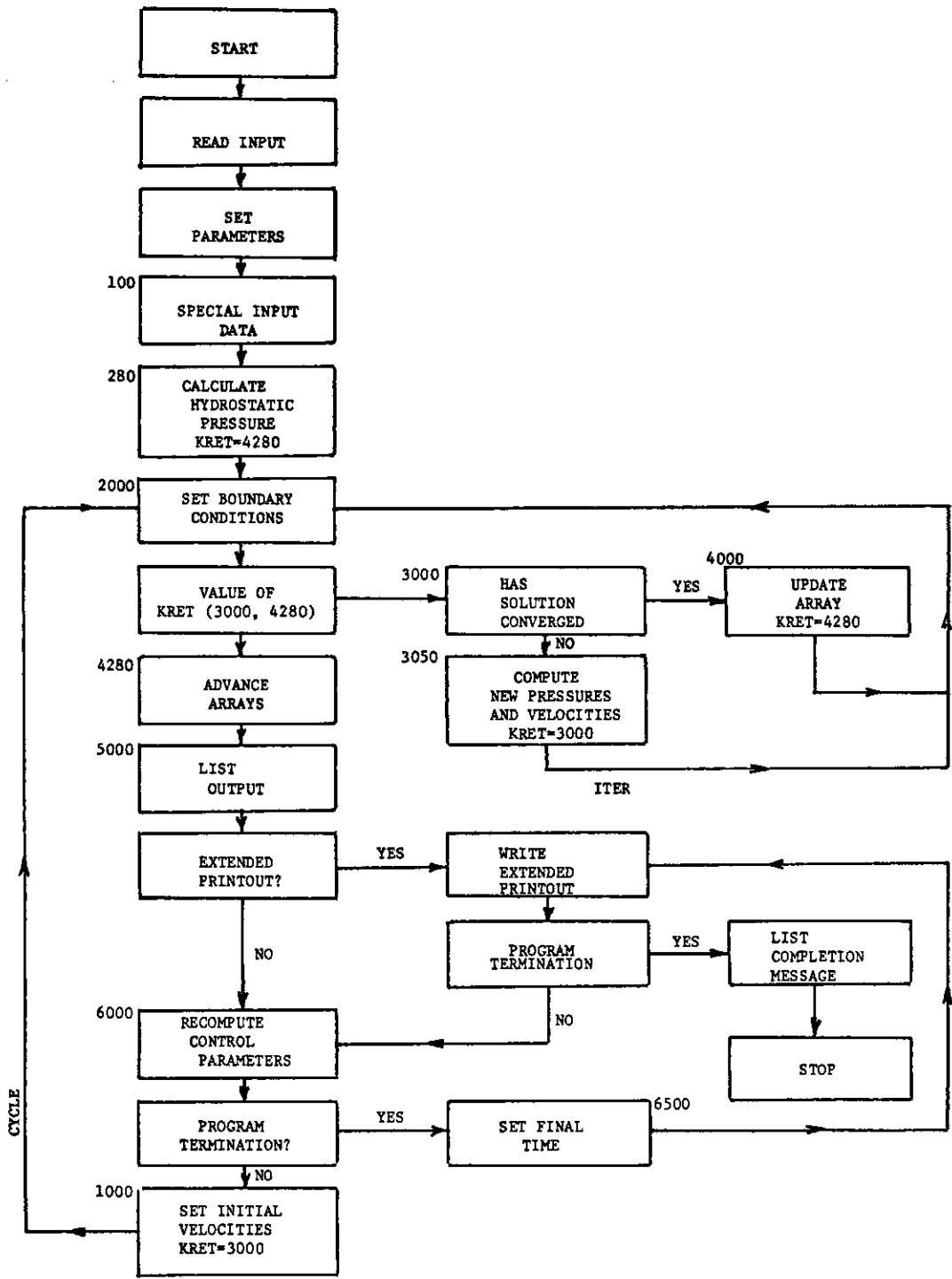


FIGURE 3: PROGRAM AVALNCH FLOW CHART

with more rectangular arrays, than with AVALNCH, in order to account for vertical variations in the flow. To specify a more rectangular array, not only must IBAR and JBAR be changed, but also the array specifications in the COMMON block at the start of the program must be changed. Application of program AVALNCH (in modified form) to impact analysis is reported by Lang and Brown, 1980. A listing of program AVALNCH that operates on a MELCOM 70 Computer System (Mitsubishi Electric, Japan) is given in Table 2.

Program AVALNCH has been used to model not only snow avalanches, but also other natural phenomena that involve transient fluid dynamics. For example, the mud flows associated with the 1980 eruption of Mt. St. Helens have been evaluated using the program (Lang, Dent, 1983). Also, a large volume rockslide that occurred in southwestern Montana following an earthquake in 1959, has been successfully modeled (Trunk, Dent, Lang, 1983).

TABLE 2: Listing of Program AVALNCH for operation on a MELCOM 70 Computer System.

```

LINE.      1      2      3      4      5      6      7
1234567890123456789012345678901234567890123456789012345678901
1          C  *  *  *  *  PROGRAM AVALNCH  *  *  *  *
2          DIMENSION U(202,4),V(202,4),UN(202,4),VN(202,4),P(202,4),
3          IH(202),HN(202),FR(202),JT(202),GX(202),GY(202),NAME(40)
4          READ(7,1) NAME
5          WRITE(6,2)
6          WRITE(6,3) NAME
7          C  *  *  *  *  READ INITIAL DATA  *  *  *  *
8          READ(7,19) IBAR, JBAR, DX, DY, YU, FK, TF, NP
9          WRITE(6,5) IBAR, JBAR, DX, DY, YU, FK, TF, NP
10         1 FORMAT(40A2)
11         2 FORMAT(1H1)
12         3 FORMAT(5X,40A2)
13         4 FORMAT(8F10.0)
14         5 FORMAT(1H ,1X,'IBAR=',I4,2X,'JBAR=',I3,2X,'DX=',F6,2,2X,'DY=',
15         1F5,2,2X,'YU=',F4,2,2X,'FK=',F4,2,2X,'TF=',F5,0,2X,'NP=',I4)
16         6 FORMAT(1H0,35X,'FLOW HEIGHT')
17         7 FORMAT(8F10.3)
18         8 FORMAT(1H0,25X,'ELEVATION CHANGE FOR EACH CELL')
19         9 FORMAT(1H0,25X,'BOUNDARY FRICTION COEFFICIENTS')
20        10 FORMAT(1H0,25X,'SLOPE-PARALLEL GRAVITY COMPONENTS')
21        11 FORMAT(1H0,25X,'SLOPE-NORMAL GRAVITY COMPONENTS')
22        12 FORMAT(1H0,30X,'END OF INPUT DATA')
23        13 FORMAT(2X,'CYCLE=',I4,2X,'ITER=',I3,2X,'DELT=',1PE9,2,2X,
24        1' TIME=',E9,2,2X,'FVOL=',E9,2,2X,'UMAX=',E9,2,2X,'UEDG=',
25        2E9,2,2X,'LDEG=',I3)
26        14 FORMAT(6X,'I',7X,'J',8X,'U',13X,'V',13X,'P',13X,'H',9X,
27        1'SUR CELL')
28        15 FORMAT(4X,I3,5X,I3,4(4X,1PE10.3),6X,I2)
29        16 FORMAT(5X,'PROBLEM RUNNING TIME EXCEEDED-CALCULATIONS STOPPED')
30        17 FORMAT(5X,'AVALANCHE AT END OF GRID-CALCULATIONS TERMINATED')
31        18 FORMAT(5X,'FLOW VELOCITY NEGLIGIBLE-CALCULATIONS TERMINATED')
32        19 FORMAT(2I10,5F10.0,I10)
33         C  *  *  *  SET PARAMETERS  *  *
34         IMAX=IBAR+2
35         JMAX=JBAR+2
36         RDX=1.0/DX
37         RDY=1.0/DY
38         IM1=IBAR+1
39         JM1=JBAR+1
40         DM=DY/100.
41         DT=1.0
42         T=0.0
43         FLG=0.0
44         UEDG1=0.0
45         AB=1.0
46         NC=0
47         ITER=0
48         IND=0
49         LDEG=0
50         G=9.806
    
```

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LINE.	1	2	3	4	5	6	7
	123456789012345678901234567890123456789012345678901234567890123456789012345678901						
51		OMG=1.7					
52		EPSI=.001					
53		ALPHA=0.1					
54		GAMMA=0.1					
55		DZRO=1.0					
56		BETA=OMG/(2.*DT*(RDX**2+RDY**2))					
57		IF(NF.EQ.1) ICP=2					
58		DO 100 I=1,IMAX					
59		H(I)=0.0					
60		HN(I)=0.0					
61		FR(I)=0.0					
62		JT(I)=0					
63		GX(I)=0.0					
64		GY(I)=0.0					
65		DO 100 J=1,JMAX					
66		U(I,J)=0.0					
67		V(I,J)=0.0					
68		UN(I,J)=0.0					
69		VN(I,J)=0.0					
70		P(I,J)=0.0					
71		100 CONTINUE					
72	C	* * * * SPECIAL INPUT DATA * * * *					
73	C						
74	C	* * FLOW HEIGHT * *					
75		READ(7,4)(H(I),I=2,IM1)					
76		WRITE(6,6)					
77		WRITE(6,7)(H(I),I=2,IM1)					
78	C	* * ELEVATION CHANGE FOR EACH CELL * *					
79		READ(7,4)(HN(I),I=2,IM1)					
80		WRITE(6,8)					
81		WRITE(6,7)(HN(I),I=2,IM1)					
82	C	* * FRICTION COEFFICIENTS * *					
83		IF(FK.GT.0.0) GO TO 120					
84		READ(7,4)(FR(I),I=2,IM1)					
85		GO TO 130					
86	120	DO 125 I=2,IM1					
87	125	FR(I)=FK					
88	130	CONTINUE					
89		WRITE(6,9)					
90		WRITE(6,7)(FR(I),I=2,IM1)					
91	C	* * GRAVITY COMPONENTS * *					
92		DO 150 I=2,IM1					
93		SI=HN(I)/DX					
94		CO=SQRT(1.0-SI*SI)					
95		GX(I)=G*SI					
96	150	GY(I)=-G*CO					
97		WRITE(6,10)					
98		WRITE(6,7)(GX(I),I=2,IM1)					
99		WRITE(6,11)					
100		WRITE(6,7)(GY(I),I=2,IM1)					

```

-----
LINE.      1      2      3      4      5      6      7
12345678901234567890123456789012345678901234567890123456789012345678901
101      C * * END OF INPUT DATA * *
102      WRITE(6,12)
103      C * * SET CELL NUMBER OF FLOW HEIGHT * *
104      DO 240 I=2,IM1
105      JT(I)=INT(H(I)*RBY+0.001)+2
106      IF(JT(I).GT.JM1) JT(I)=JM1
107      240 HN(I)=0.0
108      H(1)=H(2)
109      H(IMAX)=H(IM1)
110      JT(1)=JT(2)
111      JT(IMAX)=JT(IM1)
112      C * * CALCULATE HYDROSTATIC PRESSURE * *
113      DO 280 I=2,IM1
114      JT1=JT(I)
115      DO 280 J=2,JT1
116      280 P(I,J)=-GY(I)*(H(I)-(FLOAT(J)-1.5)*DY)
117      ASSIGN 4280 TO KRET
118      GO TO 2000
119      C * * START CYCLE OF COMPUTATIONS * *
120      1000 CONTINUE
121      ITER=0
122      FLG=1.0
123      ASSIGN 3000 TO KRET
124      C * * COMPUTE TEMPORARY U AND V VELOCITIES * *
125      DO 1100 I=2,IM1
126      JT1=JT(I)
127      DO 1100 J=2,JT1
128      FUX=((UN(I,J)+UN(I+1,J))*(UN(I,J)+UN(I+1,J))+ALPHA*ABS(UN(I,J)
129      1+UN(I+1,J))*(UN(I,J)-UN(I+1,J))-UN(I-1,J)+UN(I,J))*(UN(I-1,J)
130      2+UN(I,J))-ALPHA*ABS(UN(I-1,J)+UN(I,J))*(UN(I-1,J)-UN(I,J)))/(4
131      3.0*DX)
132      FUY=((VN(I,J)+VN(I+1,J))*(UN(I,J)+UN(I,J+1))+ALPHA*ABS(VN(I,J)
133      1+VN(I+1,J))*(UN(I,J)-UN(I,J+1))-VN(I,J-1)+VN(I+1,J-1))*(UN(I,
134      2J-1)+UN(I,J))-ALPHA*ABS(VN(I,J-1)+VN(I+1,J-1))*(UN(I,J-1)-UN(I
135      3,J)))/(4.0*DY)
136      FVX=((UN(I,J)+UN(I,J+1))*(VN(I,J)+VN(I+1,J))+ALPHA*ABS(UN(I,J)
137      1+UN(I,J+1))*(VN(I,J)-VN(I+1,J))-UN(I-1,J)+UN(I-1,J+1))*(VN(I-
138      21,J)+VN(I,J))-ALPHA*ABS(UN(I-1,J)+UN(I-1,J+1))*(VN(I-1,J)-VN(I
139      3,J)))/(4.0*DX)
140      FVY=((VN(I,J)+VN(I,J+1))*(VN(I,J)+VN(I,J+1))+ALPHA*ABS(VN(I,J)
141      1+VN(I,J+1))*(VN(I,J)-VN(I,J+1))-VN(I,J-1)+VN(I,J-1))*(VN(I,J-1)
142      2+VN(I,J))-ALPHA*ABS(VN(I,J-1)+VN(I,J))*(VN(I,J-1)-VN(I,J)))/(4
143      3.0*DY)
144      VISX=YU*((UN(I+1,J)-2.*UN(I,J)+UN(I-1,J))/DX**2+(UN(I,J+1)-2.*
145      1UN(I,J)+UN(I,J-1))/DY**2)
146      VISY=YU*((VN(I+1,J)-2.*VN(I,J)+VN(I-1,J))/DX**2+(VN(I,J+1)-2.*V
147      1N(I,J)+VN(I,J-1))/DY**2)
148      U(I,J)=UN(I,J)+DT*((P(I,J)-P(I+1,J))*RDX+GX(I)-FUX-FUY+VISX)
149      1100 V(I,J)=VN(I,J)+DT*((P(I,J)-P(I,J+1))*RBY+GY(I)-FVX-FVY+VISY)
150      C * * SET BOUNDARY CONDITIONS * *
-----

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LINE.	1	2	3	4	5	6	7
151	2000	CONTINUE					
152		HN(1)=HN(2)					
153		HN(IMAX)=HN(IM1)					
154		JT(1)=JT(2)					
155		JT(IMAX)=JT(IM1)					
156	C	*	*	LEFT WALL RIGID AND SLIP FREE	*	*	
157	C	*	*	RIGHT WALL CONTINUOUS OUTFLOW	*	*	
158		DO 2200 J=1, JMAX					
159		U(1, J)=0.0					
160		V(1, J)=V(2, J)					
161		IF(ITER.GT.0) GO TO 2200					
162		U(IM1, J)=U(IBAR, J)					
163		2200 V(IMAX, J)=V(IM1, J)					
164	C	*	*	TOP WALL CONTINUOUS OUTFLOW	*	*	
165	C	*	*	BOTTOM WALL RIGID WITH FRICTION	*	*	
166		DO 2500 I=1, IMAX					
167		IF(ITER.GT.0) GO TO 2400					
168		V(I, JM1)=V(I, JBAR)					
169		U(I, JMAX)=U(I, JM1)					
170		2400 V(I, 1)=0.0					
171		2500 U(I, 1)=U(I, 2)*(1.0-2.0*FR(I))					
172	C	*	*	FREE SURFACE BOUNDARY CONDITIONS	*	*	
173		DO 2700 I=2, IM1					
174		JT1=JT(I)					
175		IF(JT(I+1).LT.JT(I)) U(I, JT1)=U(I, JT1-1)					
176		V(I, JT1)=V(I, JT1-1)-DY*RDY*(U(I, JT1)-U(I-1, JT1))					
177		2700 U(I, JT1+1)=U(I, JT1)					
178		GO TO KRET, (3000, 4280)					
179		3000 CONTINUE					
180	C	*	*	CHECK FOR CONVERGENCE	*	*	
181		IF(FLG.EQ.0.) GO TO 4000					
182		ITER=ITER+1					
183		IF(ITER.LT.500) GO TO 3050					
184		T=1.E+10					
185		GO TO 4000					
186		3050 FLG=0.0					
187	C	*	*	COMPUTE UPDATED CELL PRESSURE AND VELOCITIES	*	*	
188		JB1=2					
189		DO 3500 I=2, IM1					
190		JT1=JT(I)					
191		DO 3500 J=2, JT1					
192		IF(JT1.EQ.JB1) GO TO 3060					
193		IF(J.NE.JB1.AND.J.NE.JT1) GO TO 3200					
194		IF(J.EQ.JT1) GO TO 3100					
195		GO TO 3200					
196		3060 CONTINUE					
197		F=V(I, J)+DY*RDY*(U(I, J)-U(I-1, J))					
198		DFDP=DT*RDY*(1.0+2.0*DY**2+RDY**2)					
199		DF1=-F/DFDP					
200		3100 ETA=DY/(HN(I)-(FLOAT(JT1)-2.5)*DY)					

LINE.	1	2	3	4	5	6	7
201							
202							
203	3200						
204							
205							
206	3300						
207							
208							
209							
210	3500						
211							
212	4000						
213	C *						
214							
215							
216							
217							
218							
219							
220							
221							
222	4100						
223	C *						
224	C *						
225							
226							
227							
228							
229	4250						
230							
231							
232	4280						
233	C *						
234							
235							
236	4300						
237							
238	C *						
239							
240							
241	4400						
242							
243							
244	4500						
245							
246	4600						
247							
248							
249	4700						
250	C *						

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LINE.	1	2	3	4	5	6	7
	12345678901234567890123456789012345678901234567890123456789012345678901						
251							IF(LDEG.EQ.LDEG1) GO TO 4800
252							IF(NC.GT.0) UEDG=DX/TC
253							IF(NC.EQ.0) UEDG=5.0
254							TC=DT
255							INFLO=1
256							IF(UEDG.GT.UEDG1) UEDG1=UEDG
257							GO TO 4910
258	4800						TC=TC+DT
259							INFLO=INFLO+1
260	C	*	*	ADVANCE U,V,H ARRAYS	*	*	
261	4910						UM=0.0
262							VM=0.0
263							DO 4900 I=1,IMAX
264							DO 4900 J=1,JMAX
265							UA=ABS(U(I,J))
266							VA=ABS(V(I,J))
267							PA=ABS(P(I,J))
268							IF(UA.GT.1.0E+04) U(I,J)=0.0
269							UN(I,J)=U(I,J)
270							IF(VA.GT.1.0E+04) V(I,J)=0.0
271							VN(I,J)=V(I,J)
272							IF(PA.LT.1.0E-16) P(I,J)=0.0
273	4900						HN(I)=H(I)
274							DO 4950 I=KTEG,LDEG
275							DO 4950 J=2,JM1
276							UT=ABS(UN(I,J))
277							VT=ABS(VN(I,J))
278							IF(UT.GT.UM) UM=UT
279	4950						IF(VT.GT.VM) VM=VT
280	C	*	*	LIST VELOCITY,PRESSURE,AND SURFACE POSITION	*	*	
281	5000			WRITE(6,13)NC,ITER,DT,T,FV,UM,UEDG,LDEG			
282							IF(NC.EQ.ICP) GO TO 5030
283							GO TO 6000
284	5030			ICP=ICP+NP			
285	5060			CONTINUE			
286				WRITE(6,14)			
287				DO 5250 I=1,IMAX			
288				JT1=JT(I)			
289				JT2=JT1+1			
290				DO 5250 J=1,JT2			
291				WRITE(6,15)I,J,U(I,J),V(I,J),P(I,J),H(I),JT1			
292	5250			CONTINUE			
293				GO TO (6000,6520,6530,6540),IND			
294	C	*	*	RECOMPUTE CONTROL PARAMETERS	*	*	
295	6000			IF(NC.EQ.0) GO TO 6300			
296				DTX=DX/UM			
297				DTY=DY/VM			
298				DT=AMIN1(DTX,DTY)/3.0			
299				IF(ITER.LT.10) DT=1.5*DT			
300				YU1=YU-1.0E-06			

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LINE.      1      2      3      4      5      6      7
1234567890123456789012345678901234567890123456789012345678901
301          IF(YU1.LT.0.0) GO TO 6300
302          DET=(DX*DY)**2/(2.*YU*(DX**2+DY**2))
303          IF(DT.LT.DET) GO TO 6300
304          DT=0.9*DET
305      6300  T=T+DT
306          IF(NC.EQ.0) GO TO 6400
307          DAX=UM*DT/DX
308          DAY=VM*DT/DY
309          ALPHA=1.35*AMAX1(DAX, DAY)
310          IF(ALPHA.GT.1.0) ALPHA=0.95
311          GAMMA=ALPHA
312          BETA=OMG/(2.*DT*(RDX**2+RDY**2))
313      C    *    *    TEST FOR PROGRAM TERMINATION    *    *
314      6400  IF(T.GT.TF) IND=2
315          IF(H(IBAR).GT.DM) IND=3
316          IF(INFLO.EQ.50) IND=4
317          UEDG2=0.05*UEDG1
318          IF(UED3.LT.UEDG2) IND=4
319          IF(IND.GT.1) GO TO 6500
320          IF(NC.LT.3) GO TO 6440
321          AA=1.0+20.0*EXP(-1.25*UEDG)
322          DO 6430 I=2, IM1
323      6430  FR(I)=FR(I)*AA/AB
324          AB=AA
325      6440  NC=NC+1
326          GO TO 1000
327      6500  T=T-DT
328          GO TO 5060
329      6520  WRITE(6,16)
330          GO TO 6600
331      6530  WRITE(6,17)
332          GO TO 6600
333      6540  WRITE(6,18)
334      6600  STOP
335          END

```

*** PRINT END ***

PROGRAM AVALNCH WITH BI-VISCOUS OPTION

From observations of avalanche flows it has long been recognized that avalanches decelerate at increasing rates as they come to a stop. Thus, disaggregated snow has the general fluid property of thixotropy, and in order to numerically model avalanche flow it is necessary to account for this thixotropic condition in some way. In program AVALNCH this was accomplished by increasing the friction coefficient, f , as the flow slowed down. To do this it was necessary to prescribe a speed below which the increase starts, and a rate of increase of f as the speed continues to decrease. From observations by Schaerer (1975) on a number of avalanches at Rodgers' Pass, Canada, a transition in the speed range $5-10\text{ms}^{-1}$ was noted, so a transition speed of 5.0ms^{-1} was used in the program. No data existed from which the rate of increase of f with decrease in speed could be established. So numerical experimentation was carried out using different rates, until the stopping distances of several avalanches were matched with site measurements. It was determined that a geometric-progression type increase in f as speed decreased below the transition speed of 5ms^{-1} was needed. The resulting expression for f that is used in program AVALNCH is

$$f = f_0(1+20e^{-1.25U})$$

where f_0 is the nominal high speed value of the friction coefficient. In program AVALNCH this mechanism is referred to as the "fast-stop" option.

Recognizing, physically, that the surface friction is unlikely to increase in a geometric progression with decrease in speed, a more rational expression of the fast-stop or material locking property was sought. The physical process of fluid locking is known as the Bingham fluid property, or that the fluid is a "Bingham material". With regard to snow, the effort to apply the Bingham material concept was two-fold.

One effort was to computer model small volume snow flows that had been run experimentally, in order to determine the motion of decelerating snow (Dent, Lang 1982). Results from this work showed that a biviscous rather than a Bingham representation of the snow flow best fit the experimental data. The basic difference between the two mechanisms is that the Bingham fluid has infinite viscosity below the cutoff shear stress, τ_0 , whereas the biviscous model has finite, but larger viscosity below the cutoff shear stress (Figure 4). In computer modeling the controlled volume snow flows kinematic viscosities of $\nu = 0.002\text{m}^2\text{s}^{-1}$ and $\nu' = 0.10\text{m}^2\text{s}^{-1}$ gave good correspondance between the data.

A second effort was to independently measure the cutoff shear stress, τ_0 , that separates the two regions of the biviscous model. Results from the controlled volume snow flow tests gave $\frac{\tau_0}{\rho} = 2.2\text{m}^2\text{s}^{-2}$. Simple laboratory tests on similar type snow gave corresponding values for $\frac{\tau_0}{\rho}$ (Lang, Dent, 1983). Although the velocities of the two types of experiments were vastly different, the velocity gradients were of the same magnitude, which attributes to the close correspondance between the results of the two experiments.

From the controlled volume snow flow results we note a factor of 50 between the values of ν and ν' . These findings indicate that it is the viscosity of the snow that changes value with speed, rather than the surface friction. So the question is raised if program AVALNCH can be modified to incorporate viscosity rather than friction change, as the flow velocity decreases, and yet retain the flow characteristics of snow.

To incorporate the biviscous option in program AVALNCH was a relatively simple process. The friction mechanism was changed from a partial-slip to a no-slip boundary condition at the lower boundary. This eliminates the need to input a value for friction. In its' place

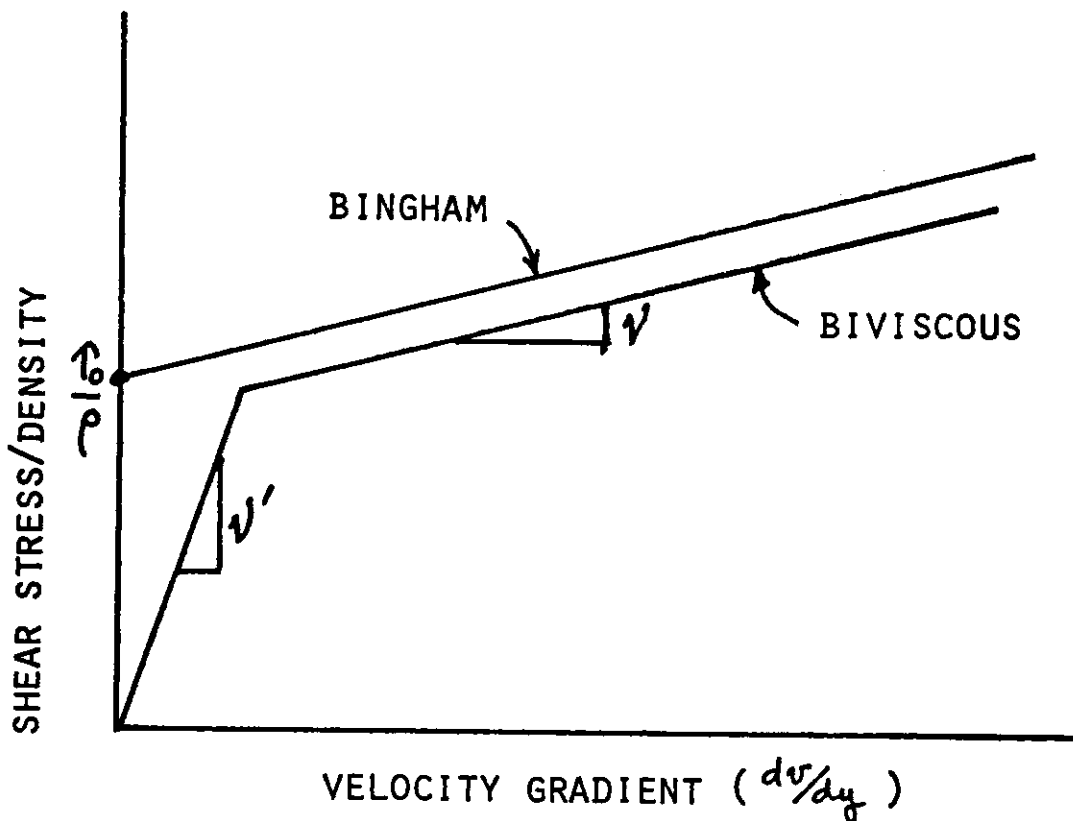


FIGURE 4: BINGHAM AND BIVISCOUS FLUID REPRESENTATION

on the input statement the multiplicative factor between ν and ν' is specified, and is designated XYU. The high stress viscosity ν is input as before, and is designated YU. The fast-stop instructions previously used were replaced by a set of instructions that test the speed of fluid in each cell along the avalanche path. If the cell has a velocity less than 5ms^{-1} , then the viscosity is set at the low stress value (XYU times YU). If the velocity is greater than 5ms^{-1} , then the viscosity is set at the high stress value (YU). These few changes were what were needed to modify the code to the biviscous option. What remained was to determine the values of ν and ν' in order to model an avalanche flow.

To evaluate the viscosities the Ironton Park avalanche was selected. In increasing the friction from $f = 0.5$ (partial-slip) to $f = 1.0$ (no-slip), the viscosity must be decreased from $\nu = 0.5\text{m}^2\text{s}^{-1}$ for the partial-slip (fast-stop) case. In addition the multiplicative factor between ν and ν' needs to be specified. Although a factor of 50 between ν and ν' was noted earlier for the small volume snow flow experiments, this factor is known not to be sensitive, so only a factor of 10 was used in the Ironton Park evaluation. From runs of Ironton Park, it was determined that $\nu = 0.23\text{m}^2\text{s}^{-1}$ yielded a duplication in runout distance and maximum velocity between the fast-stop and biviscous options. A comparison of velocities along the path is also necessary in order to establish correspondance (Figure 5), which is also achieved to a close approximation.

The corrections described for the biviscous version of AVALNCH are sufficient if the modeling is with constant values of friction and viscosity along the avalanche path. However, for some avalanches the viscosity should be variable along the path, as in the case of coastal avalanches, where part of the runout may be on dry snow, and another part on wet snow. With the no-slip boundary condition, which is the usual boundary condition in fluid dynamics, it must be assumed that

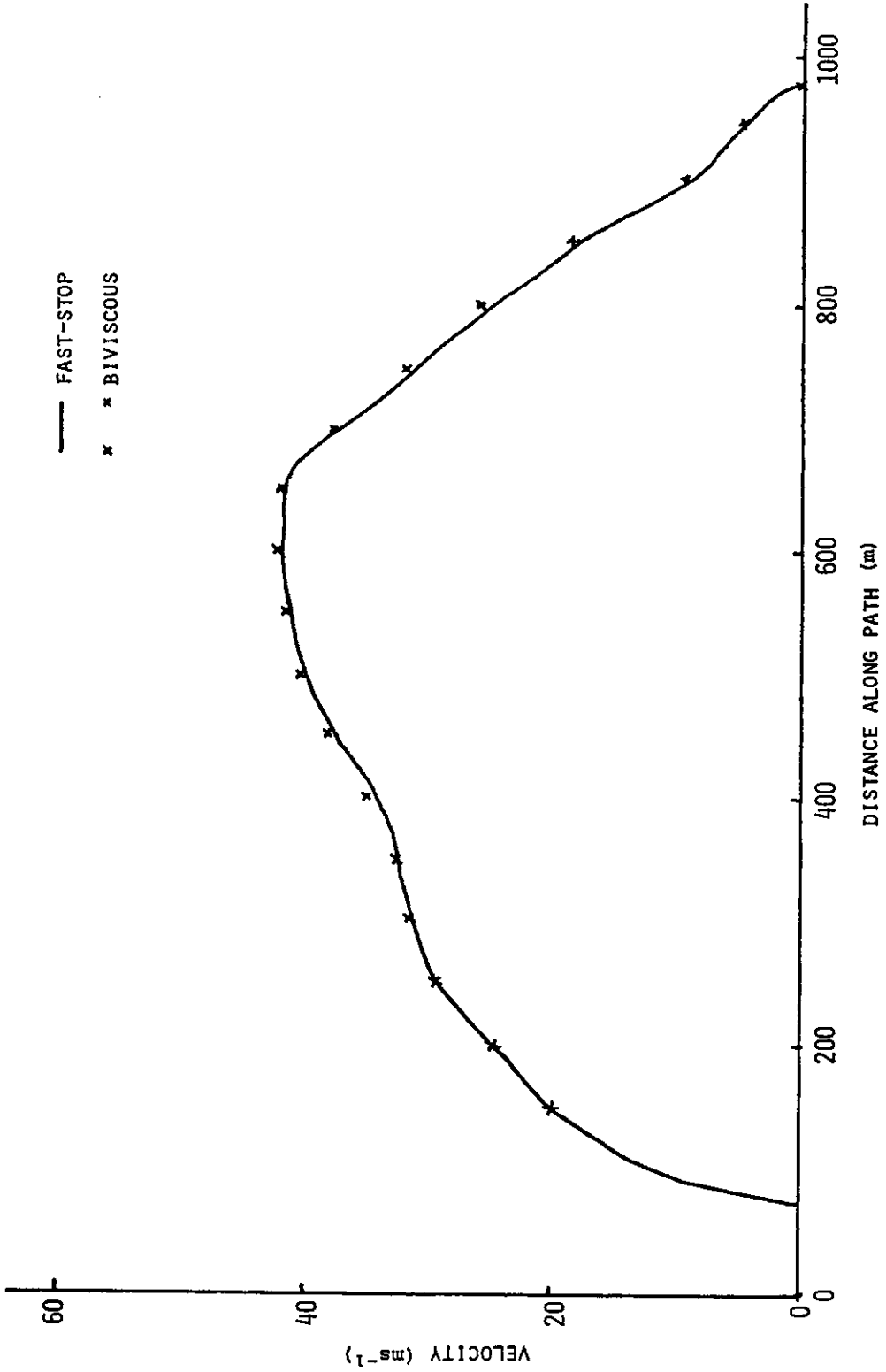


FIGURE 5: IRONTON PARK AVALANCHE VELOCITY PROFILES USING FAST-STOP AND BIVISCOUS OPTIONS.

the increased resistance of the wet snow produces additional internal mixing in the moving snow, and hence an increase in viscous dissipation. To incorporate a variable viscosity into the code it is necessary to set up viscosity as a one-dimensional array so that values of ν may be assigned to each cell along the path. This was done in a way similar to the variable friction option in the fast-stop version of the code. The change is reflected in the modified input format for the biviscous case, which is listed below.

BIVISCOUS-AVALNCH INPUT FORMAT

Input may be by a data file or by a sequence of cards depending upon user preference.

Line 1: FORMAT (40A2)

Columns 1-80: title and identification information

Line 2: FORMAT (2I10, 5F10.0, I10)

- Columns 1-10: IBAR - number of cells in the slope-parallel direction; maximum is 200, unless program is changed.
- Columns 11-20: JBAR - number of cells normal to the path; maximum is 2, unless program is changed.
- Columns 21-30: DX - dimension of cell along path (m).
- Columns 31-40: DY - dimension of cell normal to path (m).
- Columns 41-50: YU - high shear stress kinematic viscosity (m^2s^{-1}).
- Columns 51-60: XYU - multiplicative factor for low shear stress viscosity.
- Columns 61-70: TF - avalanche flow time (s).
- Columns 71-80: NP - number of cycles between extended printouts.

Line 3: FORMAT (8F10.0)

The thickness of the avalanche slab at initial release is listed in the same format as on page 9.

Line 4: FORMAT (8F10.0)

The elevation change of each cell along the avalanche path is listed in the same format as on page 9.

Line 5: FORMAT (8F10.0)

If YU=0 on Line 2, then viscosity must be specified for each cell along the path. This listing of viscosity is in the same format as for friction coefficients on page 10.

A listing of the Biviscous version of program AVALNCH is given in Table 3. The flow chart for the Biviscous version is the same as for the fast-stop version, as listed on page 11. Both the fast-stop and biviscous coefficient selection statements are in the 6000 section of the program.

TABLE 3 : Listing of Biviscous Version of AVALNCH.

LINE.	1	2	3	4	5	6	7
1	C	*	*	BIAV: PROGRAM AVALNCH WITH BIVISCOUS OPTION	*	*	
2				DIMENSION U(202,4),V(202,4),UN(202,4),VN(202,4),P(202,4),			
3				1H(202),HN(202),XU(202,2),JT(202),GX(202),GY(202),NAME(40)			
4				READ(7,1) NAME			
5				WRITE(6,2)			
6				WRITE(6,3) NAME			
7	C	*	*	* * * * READ INITIAL DATA	*	*	*
8				READ(7,19) IBAR,JBAR,DX,DY,YU,XYU,TF,NP			
9				WRITE(6,5) IBAR,JBAR,DX,DY,YU,XYU,TF,NP			
10				READ(7,4) DTZ,UZ			
11				WRITE(6,20) DTZ,UZ			
12				1 FORMAT(40A2)			
13				2 FORMAT(1H1)			
14				3 FORMAT(5X,40A2)			
15				4 FORMAT(8F10.0)			
16				5 FORMAT(1H,1X,'IBAR=',I4,2X,'JBAR=',I3,2X,'DX=',F6,2,2X,'DY=',			
17				1F5,2,2X,'YU=',F4,2,2X,'XYU=',F4,1,2X,'TF=',F5,0,2X,'NP=',I4)			
18				6 FORMAT(1H0,35X,'FLOW HEIGHT')			
19				7 FORMAT(8F10.3)			
20				8 FORMAT(1H0,25X,'ELEVATION CHANGE FOR EACH CELL')			
21				9 FORMAT(1H0,25X,'VISCOSITY VALUES FOR EACH CELL')			
22				10 FORMAT(1H0,25X,'SLOPE-PARALLEL GRAVITY COMPONENTS')			
23				11 FORMAT(1H0,25X,'SLOPE-NORMAL GRAVITY COMPONENTS')			
24				12 FORMAT(1H0,30X,'END OF INPUT DATA')			
25				13 FORMAT(2X,'CYCLE=',I4,2X,'ITER=',I3,2X,'DELT=',1PE9,2,2X,			
26				1'TIME=',E9,2,2X,'FVOL=',E9,2,2X,'UMAX=',E9,2,2X,'UEDG=',			
27				2E9,2,2X,'LDEG=',I3)			
28				14 FORMAT(6X,'I',7X,'J',8X,'U',13X,'V',13X,'P',13X,'H',9X,			
29				1'SUR CELL')			
30				15 FORMAT(4X,I3,5X,I3,4(4X,1PE10.3),6X,I2)			
31				16 FORMAT(5X,'PROBLEM RUNNING TIME EXCEEDED-CALCULATIONS STOPPED')			
32				17 FORMAT(5X,'AVALANCHE AT END OF GRID-CALCULATIONS TERMINATED')			
33				18 FORMAT(5X,'FLOW VELOCITY NEGLIGIBLE-CALCULATIONS TERMINATED')			
34				19 FORMAT(2I10,5F10.0,I10)			
35				20 FORMAT(5X,'DTZ=',F6,3,5X,'UZ=',F6,3)			
36	C	*	*	SET PARAMETERS	*	*	
37				IMAX=IBAR+2			
38				JMAX=JBAR+2			
39				RDX=1.0/DX			
40				RDY=1.0/DY			
41				IM1=IBAR+1			
42				JM1=JBAR+1			
43				DM=DY/100.			
44				DT=DTZ			
45				T=0.0			
46				FLG=0.0			
47				UEDG1=0.0			
48				ICP=NP			
49				NC=0			
50				ITER=0			

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LINE.	1	2	3	4	5	6	7
51							
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74							
75							
76	C	*	*	*	*	SPECIAL INPUT DATA	*
77	C						
78	C	*	*	FLOW HEIGHT	*	*	
79				READ(7,4)	(H(I),I=2,IM1)		
80				WRITE(6,6)			
81				WRITE(6,7)	(H(I),I=2,IM1)		
82	C	*	*	ELEVATION CHANGE FOR EACH CELL	*	*	
83				READ(7,4)	(HN(I),I=2,IM1)		
84				WRITE(6,8)			
85				WRITE(6,7)	(HN(I),I=2,IM1)		
86	C	*	*	INITIALIZE VISCOSITY	*	*	
87				IF(YU.EQ.0.0)	GO TO 125		
88				DO 120	I=2,IM1		
89				XU(I,1)=YU			
90				120 XU(I,2)=YU			
91				GO TO 130			
92				125 READ(7,4)	(XU(I,1),I=2,IM1)		
93				DO 127	I=2,IM1		
94				127 XU(I,2)=XU(I,1)			
95				WRITE(6,9)			
96				WRITE(6,7)	(XU(I,1),I=2,IM1)		
97				130 CONTINUE			
98	C	*	*	GRAVITY COMPONENTS	*	*	
99				DO 150	I=2,IM1		
100				SI=HN(I)/DX			

LINE.	1	2	3	4	5	6	7
101							
102							
103	150						
104							
105							
106							
107							
108	C	*	*	*	*	*	*
109							
110	C	*	*	*	*	*	*
111							
112							
113							
114	240						
115							
116							
117							
118							
119	C	*	*	*	*	*	*
120							
121							
122							
123	280						
124							
125							
126	C	*	*	*	*	*	*
127	1000						
128							
129							
130							
131	C	*	*	*	*	*	*
132							
133							
134							
135							
136							
137							
138							
139							
140							
141							
142							
143							
144							
145							
146							
147							
148							
149							
150							

Computer Programs for Avalanche Runout Prediction—Lang

LINE.	1	2	3	4	5	6	7
151	VISX=XU(I,1)*((UN(I+1,J)-2.*UN(I,J)+UN(I-1,J))/DX**2+(UN(I,J+1)						
152	1-2.*UN(I,J)+UN(I,J-1))/DY**2)						
153	VISY=XU(I,1)*((VN(I+1,J)-2.*VN(I,J)+VN(I-1,J))/DX**2+(VN(I,J+1)						
154	1-2.*VN(I,J)+VN(I,J-1))/DY**2)						
155	U(I,J)=UN(I,J)+DT*((P(I,J)-P(I+1,J))*RDX+GX(I)-FUX-FUY+VISX)						
156	1100	V(I,J)=VN(I,J)+DT*((P(I,J)-P(I,J+1))*RDY+GY(I)-FVX-FVY+VISY)					
157	C	*	*	SET	BOUNDARY	CONDITIONS	*
158	2000 CONTINUE						
159	HN(1)=HN(2)						
160	HN(IMAX)=HN(IM1)						
161	JT(1)=JT(2)						
162	JT(IMAX)=JT(IM1)						
163	C	*	*	LEFT	WALL	RIGID	AND
164	C	*	*	RIGHT	WALL	CONTINUOUS	OUTFLOW
165	DO 2200 J=1,JMAX						
166	U(1,J)=0.0						
167	V(1,J)=V(2,J)						
168	IF(ITER.GT.0) GO TO 2200						
169	U(IM1,J)=U(IBAR,J)						
170	2200	V(IMAX,J)=V(IM1,J)					
171	C	*	*	TOP	WALL	CONTINUOUS	OUTFLOW
172	C	*	*	BOTTOM	WALL	RIGID	WITH
173	DO 2500 I=1,IMAX						
174	IF(ITER.GT.0) GO TO 2400						
175	V(I,JM1)=V(I,JBAR)						
176	U(I,JMAX)=U(I,JM1)						
177	2400	V(I,1)=0.0					
178	2500	U(I,1)=-U(I,2)					
179	C	*	*	FREE	SURFACE	BOUNDARY	CONDITIONS
180	DO 2700 I=2,IM1						
181	JT1=JT(I)						
182	IF(JT(I+1).LT.JT(I)) U(I,JT1)=U(I,JT1-1)						
183	V(I,JT1)=V(I,JT1-1)-DY*RDX*(U(I,JT1)-U(I-1,JT1))						
184	2700	U(I,JT1+1)=U(I,JT1)					
185	GO TO KRET,(3000.4280)						
186	3000 CONTINUE						
187	C	*	*	CHECK	FOR	CONVERGENCE	*
188	IF(FLG.EQ.0.) GO TO 4000						
189	ITER=ITER+1						
190	IF(ITER.LT.500) GO TO 3050						
191	T=1.E+10						
192	GO TO 4000						
193	3050	FLG=0.0					
194	C	*	*	COMPUTE	UPDATED	CELL	PRESSURE
195	AND						
196	VELOCITIES						
197	*						
198	JB1=2						
199	DO 3500 I=2,IM1						
200	JT1=JT(I)						
	DO 3500 J=2,JT1						
	IF(JT1.EQ.JB1) GO TO 3060						
	IF(J.NE.JB1.AND.J.NE.JT1) GO TO 3200						

LINE.	1	2	3	4	5	6	7
201							
202							
203	3060						
204							
205							
206							
207	3100						
208							
209							
210	3200						
211							
212							
213	3300						
214							
215							
216							
217	3500						
218							
219	4000						
220	C	*	*				
221							
222							
223							
224							
225							
226							
227							
228							
229	4100						
230	C	*	*				
231	C	*	*				
232							
233							
234							
235							
236	4250						
237							
238							
239	4280						
240	C	*	*				
241							
242							
243	4300						
244							
245	C	*	*				
246							
247							
248	4400						
249							
250							

Computer Programs for Avalanche Runout Prediction—Lang

LINE.	1	2	3	4	5	6	7
251	4500	LDEG=I					
252		I=2					
253	4600	IF (H(I).GT.DM) GO TO 4700					
254		I=I+1					
255		GO TO 4600					
256	4700	KTEG=I					
257	C *	* ADVANCE U,V,H ARRAYS * *					
258	4910	UM=0.0					
259		VM=0.0					
260		DO 4900 I=1,IMAX					
261		DO 4900 J=1,JMAX					
262		UA=ABS(U(I,J))					
263		VA=ABS(V(I,J))					
264		PA=ABS(P(I,J))					
265		IF (UA.GT.1.0E+04) U(I,J)=0.0					
266		UN(I,J)=U(I,J)					
267		IF (VA.GT.1.0E+04) V(I,J)=0.0					
268		VN(I,J)=V(I,J)					
269		IF (PA.LT.1.0E-16) P(I,J)=0.0					
270	4900	HN(I)=H(I)					
271		DO 4950 I=KTEG,LDEG					
272		DO 4950 J=2,JM1					
273		UT=ABS(UN(I,J))					
274		VT=ABS(VN(I,J))					
275		IF (UT.GT.UM) UM=UT					
276	4950	IF (VT.GT.VM) VM=VT					
277	C *	* COMPUTE LEADING EDGE VELOCITY * *					
278		IF (LDEG.EQ.LDEG1) GO TO 4800					
279		IF (NC.GT.0) UEDG=DX/TC					
280		IF (NC.LT.10) UEDG=UM					
281		IF (NC.EQ.0) UEDG=UZ					
282		TC=DT					
283		INFLO=1					
284		IF (UEDG.GT.UEDG1) UEDG1=UEDG					
285		GO TO 4810					
286	4800	TC=TC+DT					
287		INFLO=INFLO+1					
288	4810	CONTINUE					
289	C *	* LIST VELOCITY,PRESSURE,AND SURFACE POSITION * *					
290	5000	WRITE(6,13)NC,ITER,DT,T,FV,UM,UEDG,LDEG					
291		IF (NC.EQ.ICP) GO TO 5030					
292		GO TO 6000					
293	5030	ICP=ICP+NP					
294	5060	CONTINUE					
295		WRITE(6,14)					
296		DO 5250 I=1,IMAX					
297		JT1=JT(I)					
298		JT2=JT1+1					
299		DO 5250 J=1,JT2					
300		WRITE(6,15)I,J,U(I,J),V(I,J),P(I,J),H(I),JT1					

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LINE.	1	2	3	4	5	6	7
301	5250	CONTINUE					
302		GO TO (6000,6520,6530,6540),IND					
303	C	*	*	RECOMPUTE CONTROL PARAMETERS	*	*	
304	6000	IF(NC.EQ.0) GO TO 6300					
305		DTX=DX/UM					
306		DTY=DY/VM					
307		DT=AMIN1(DTX,DTY)/3.0					
308		IF(ITER.LT.10) DT=1.5*DT					
309		YU=XU(LDEG,1)					
310		YU1=YU-1.0E-06					
311		IF(YU1.LT.0.0) GO TO 6300					
312		DET=(DX*DY)**2/(2.*YU*(DX**2+DY**2))					
313		IF(DT.LT.DET) GO TO 6300					
314		DT=0.9*DET					
315	6300	T=T+DT					
316		IF(NC.EQ.0) GO TO 6400					
317		DAX=UM*DT/DX					
318		DAY=VM*DT/DY					
319		ALPHA=1.35*AMAX1(DAX,DAY)					
320		IF(ALPHA.GT.1.0) ALPHA=0.95					
321		GAMMA=ALPHA					
322		BETA=OMG/(2.*DT*(RDX**2+RDY**2))					
323	C	*	*	TEST FOR PROGRAM TERMINATION	*	*	
324	6400	IF(T.GT.TF) IND=2					
325		IF(H(IBAR).GT.DM) IND=3					
326		IF(INFLO.EQ.50) IND=4					
327		UEDG2=0.05*UEDG1					
328		IF(UEDG.LT.UEDG2) IND=4					
329		IF(IND.GT.1) GO TO 6500					
330		IF(NC.LT.10) GO TO 6440					
331		DO 6430 I=KTEG,LDEG					
332		IF(U(I,2).LT.5.0) XU(I,1)=XYU*XU(I,2)					
333	6430	IF(U(I,2).GE.5.0) XU(I,1)=XU(I,2)					
334	6440	NC=NC+1					
335		GO TO 1000					
336	6500	T=T-DT					
337		GO TO 5060					
338	6520	WRITE(6,16)					
339		GO TO 6600					
340	6530	WRITE(6,17)					
341		GO TO 6600					
342	6540	WRITE(6,18)					
343	6600	STOP					
344		END					

*** PRINT END ***

EQUILIBRIUM FLOW MODELS WITH MATERIAL LOCKING

The equations for equilibrium flow modeling of snow avalanche runout, originally developed by Voellmy (1955), were later adapted to computer simulation by Cheng and Perla (1979). In the computer representation the avalanche path is divided into straight-line segments of varying length and integrated forms of Voellmy's acceleration equation are applied to the flow in each segment. In the computer based formulation the Voellmy acceleration equation is expressed by

$$a = g(\sin \theta - \mu \cos \theta) - \frac{D}{M} v^2$$

In this equation θ is the slope angle, μ is the coefficient of sliding friction, g is the gravitational constant (9.806ms^{-2}), and $\frac{D}{M}$ is a drag coefficient. The stated ranges on the friction and drag parameters are:

$$0.1 \leq \mu \leq 0.5$$

$$10 \leq \frac{D}{M} \leq 10^4 \quad (\text{m})$$

which must be selected based upon the site specific conditions of each avalanche path. This broad range in parameter selection has long been a difficulty in application of the equations to different types of avalanche flow. Also inherent in this formulation is an instability that occurs when the slope angle equals the assumed friction angle of the snow.

In light of recent developments pertaining to snow avalanche flow, we consider modification of the Cheng-Perla computer program by attempting to incorporate the following conditions:

1. At low shear stress values in flowing snow the snow has a tendency to lock up; a characteristic of a thixotropic fluid.
2. Flow of avalanches involves bulk flow of the major portion of the moving material riding upon a high-velocity-gradient boundary layer of granularized snow. In the boundary layer the basic mechanism of energy dissipation is by viscous effects, as is typical of fluid processes in general.

The objective in incorporating these physical effects into the equilib-

rium flow model is to reduce the variation in parameter selection for different avalanche cases, and to eliminate the instability condition of the Voellmy equation.

EQUATION FORMULATION

Using recently developed mechanics of snow flow (Dent, Lang, 1982) a model for flow incorporates a viscous boundary layer that supports the material bulk flow. We designate the depth of the bulk material by h , depth of the boundary layer by λ , and velocity of the bulk material by V (Figure 6). Forces acting on the bulk material are gravitational, viscous, and assumed frictional (Figure 7), where τ is the viscous shear stress, and A is the area of contact between the mass segment and the boundary layer. With these forces acting the equation of motion of the mass segment is:

$$\sum F = ma = mg \sin \theta - \mu N - \tau A$$

Dividing by m and setting $N=mg \cos \theta$, an acceleration equation similar to that of Voellmy is obtained.

$$a = g(\sin \theta - \mu \cos \theta) - \frac{\tau A}{m}$$

The term $\frac{\tau A}{m}$ can be rewritten $\frac{\tau}{\rho} \cdot \frac{\rho A \lambda}{m} \cdot \frac{1}{\lambda}$ where ρ is the density of the snow. But $\rho A \lambda = m$, the mass of the bulk material, so for the equation of motion we have

$$a = g(\sin \theta - \mu \cos \theta) - \frac{1}{\lambda} \frac{\tau}{\rho}$$

If we set $\frac{\tau}{\rho} = \frac{g}{\xi} V^2$ we have the Voellmy equation, where ξ is a coefficient of dynamic resistance.

If the boundary layer is assumed to be a Newtonian fluid in a laminar flow regime, then

$$\frac{\tau}{\rho} = \nu \frac{dv}{dy}$$

and the velocity gradient is approximated by $\frac{dv}{dy} = \frac{V}{\lambda}$ to obtain

$$\frac{\tau}{\rho} = \nu \frac{V}{\lambda}$$

where ν is the kinematic viscosity. In this formulation λ is the depth of the active boundary layer upon which the bulk of flowing snow

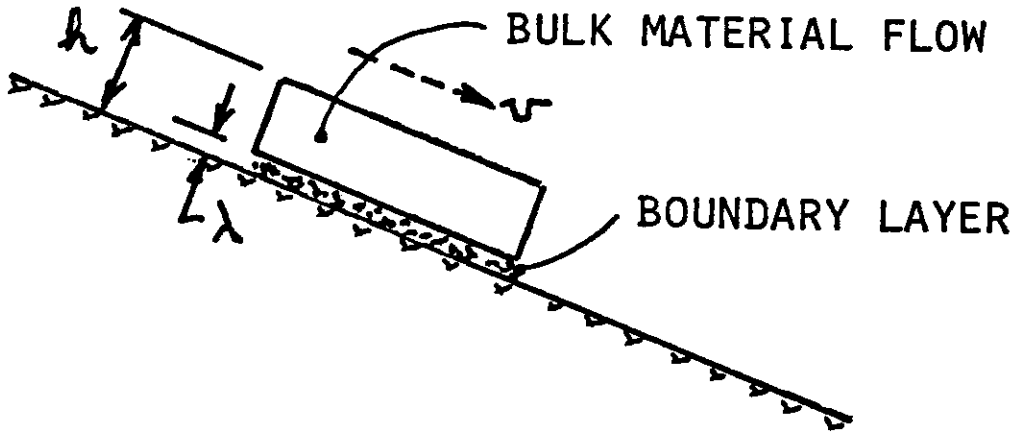


FIGURE 6: FLOWING SNOW CONFIGURATION

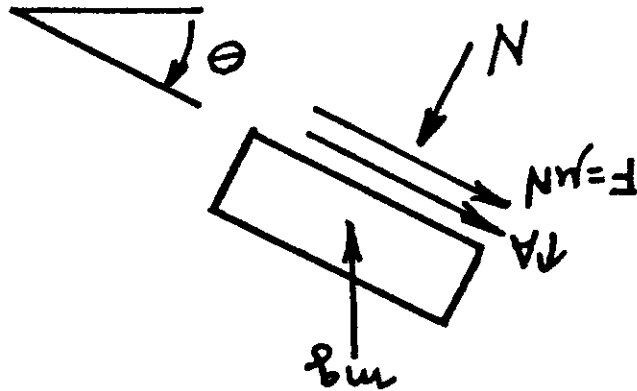


FIGURE 7: ELEMENT FORCES

rides. While the leading edge of an avalanche may exhibit strong mixing, the bulk of the dense flowing material is expected to be smooth, because of path smoothing at the leading edge. Presently, we have no basis for specifying λ , the depth of the boundary layer. In tests with small volumes of snow in controlled releases the depth observed was about 2cm for a flow depth of about 20cm. Depth of the boundary layer should depend upon the roughness over which the flow advances. For flow in pipes the depth of the boundary layer is related directly to wall roughness. In the case of avalanche flow the definition of roughness is not known. Is roughness to be related to undulations in the runout path, or to grain size in the boundary layer, or to fluxuations in the boundary layer profile? We might exclude path undulations on the basis of leading edge smoothing. However, to go beyond this to consider boundary layer profile fluxuations is more complex, and warrants experimental investigation, as has been done in the case of pipe flow. At the present time, we will express the boundary layer thickness as a fraction of the flow depth in the form

$$\lambda = r h$$

The acceleration equation under these assumptions becomes

$$a = g(\sin \theta - \mu \cos \theta) - \frac{v^*}{h^2} v$$

when $v^* = \frac{v}{h}$, is one of the basic parameters that must be evaluated.

Setting the acceleration $a = v \frac{dv}{ds}$ and integrating, we obtain the following equation relating velocity to distance of travel, s ,

$$\frac{1}{\beta^2} (\alpha - \beta v) - \frac{\alpha}{\beta^2} \ln (\alpha - \beta v) = s + C_1$$

where C_1 is a constant of integration, and

$$\alpha = g(\sin \theta - \mu \cos \theta) \quad , \quad \beta = \frac{v^*}{h^2}$$

Imposing the constraint condition that at $s=0, v=v_A$ leads to

$$s = \frac{1}{\beta^2} \left[\beta(v_A - v) + \alpha \ln \left(\frac{\alpha - \beta v_A}{\alpha - \beta v} \right) \right]$$

then, selecting at $s=L, v=v_B$, we obtain

$$\beta L = (v_A - v_B) + \frac{\alpha}{\beta} \ln \left(\frac{\alpha - \beta v_A}{\alpha - \beta v_B} \right)$$

which is a transcendental equation for velocity \bar{V}_B , assuming that in a segment analysis that \bar{V}_A would be specified from the previous segment analysis, and L is the length of the segment under evaluation. Thus, to solve for \bar{V}_B in the above equation requires the use of a numerical methods algorithm.

We note that in the laminar flow assumption that the viscous drag term in the acceleration equation is proportional to velocity to the first power. This deviates from the original Voellmy assumption, which is based upon a turbulent assumption. To investigate the turbulent flow assumption, the Boussinesq formulation (Shames, 1982) is usually cited for the relationship between shear stress in the boundary layer and the velocity gradient, namely

$$\frac{\tau}{\rho} = \nu' \left(\frac{dv}{dy} \right)^2$$

where ν' is designated the kinematic eddy viscosity, which may be related to the fluid mixing length. Making the previously defined approximations for the velocity gradient, the acceleration equation has the form

$$a = g(\sin\theta - \mu \cos\theta) - \frac{\nu^*}{h^3} v^2$$

where $\nu^* = \nu'/r^2$. Here the viscous dissipation term is proportional to v^2 , which is the same as in the Voellmy formulation. However, a difference here is that the viscous term is also proportional to h^{-3} rather than to h^{-1} , as in the Voellmy equation. Writing for the acceleration $a = \frac{1}{2} \frac{dv^2}{ds}$, integrating, and imposing the constraint conditions, as before, the following segment equations are obtained:

$$\bar{V}_B = \left[\frac{\alpha}{\beta} (1 - e^{-2\beta L}) + \bar{V}_A^2 e^{-2\beta L} \right]^{1/2}$$

Thus, in this case it is possible to express \bar{V}_B explicitly in terms of \bar{V}_A and L , the length of the segment. In a segment in which the flow stops ($\bar{V}_B = 0$) the runout distance is expressed by

$$S_f = \frac{1}{2\beta} \ln \left(1 - \frac{\beta}{\alpha} \bar{V}_A^2 \right)$$

In these equations

$$\alpha = g(\sin\theta - \mu \cos\theta) \quad , \quad \beta = \frac{v^*}{h^3}$$

so that $\frac{d}{dt}$ is not finite for the case $\alpha = 0$, which occurs if $\mu = \tan\theta$, the afore mentioned singularity condition.

MATERIAL LOCKING MECHANISM

It has long been recognized that flowing snow exhibits a thixotropic property that produces an accelerated slow-down as the avalanche comes to a stop. This locking property has been approximated by a biviscous model (Figure 8) in computer studies with multi-celled configurations (Dent, Lang, 1983). In these models the viscosity in all cells do not change simultaneously to produce a sudden change in the flow resistance. However, in the equilibrium flow models based upon Voellmy's equation only a single segment of material is used in each numerical step. Because of this simplification, it was decided that a simple biviscous approximation would be too abrupt, as the flow passed through the transition point. Instead a continuous variation in viscosity was opted for (dashed line, Figure 8), for which a functional representation is required. However, if a functional representation is selected for v^* , then the previously integrated equations are no longer applicable. Two possible recourses to this difficulty are: 1) to numerically integrate the acceleration equation in each segment analysis, or 2) to subdivide the segments into a step-wise linearized approximation. The first approach would be a radical divergence from the algorithm used by Cheng and Perla, so was discarded in favor of the second approach, which is only a modification to the Cheng-Perla program.

The next consideration is to decide upon an avalanche speed at which the viscosity transition is to occur. Data from Schaerer (1975), based upon observations of a number of avalanches at Rodger's Pass, Canada,

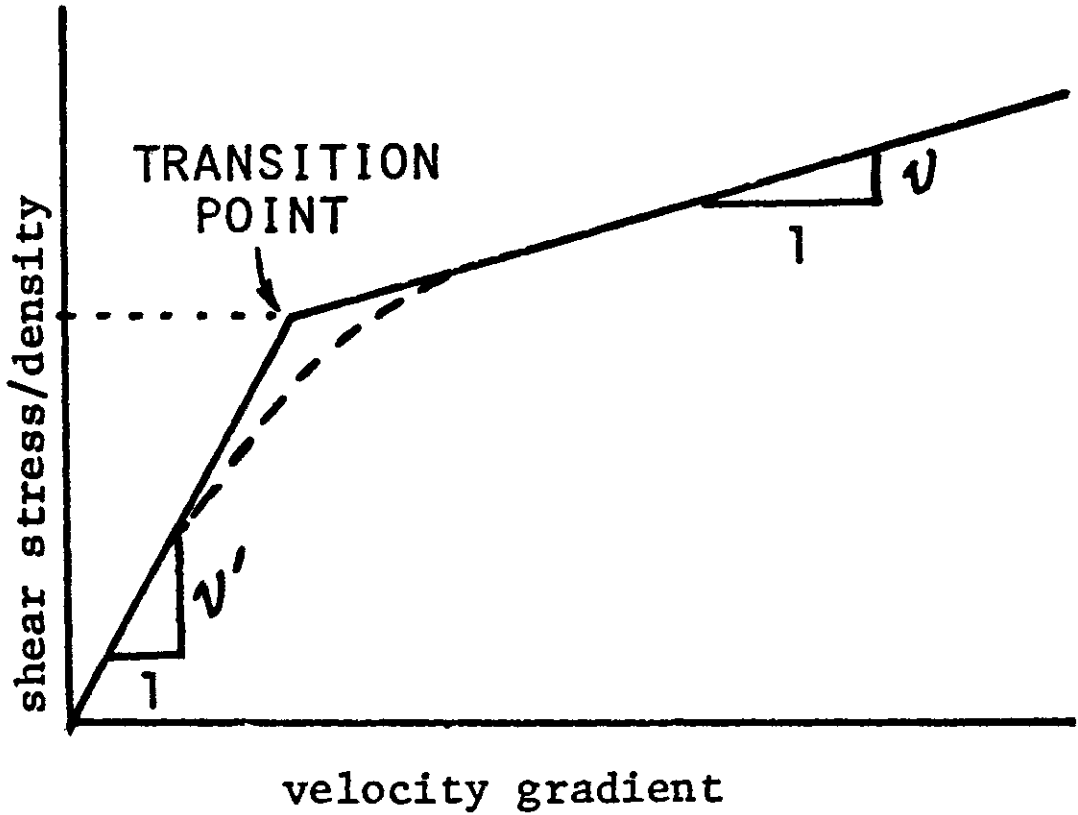


FIGURE 8: BIVISCOUS MATERIAL REPRESENTATION

place the transition in the speed range $10 > U > 5 \text{ ms}^{-1}$. This speed range has also been used with success in the fast-stop option of program AVALNCH (Lang, Dawson, Martinelli, 1979). Based upon these observations the form of the function for v^* is taken as

$$v^* = v_0 (1 + C_0 e^{-1.25 U})$$

The exponent coefficient 1.25 is sufficient for $v^* \rightarrow v_0$ for $U > 8.0 \text{ ms}^{-1}$. The coefficients v_0 and C_0 must be determined, based upon numerical study of typical avalanche flows.

The other parameter, the friction coefficient μ , in the Voellmy formulation, must be reconsidered in relation to the viscous mechanism that has now been defined. In granular material flow processes, the viscous (or velocity dependent) resistive processes are the dominant effect, and snow should be no exception. In the Voellmy formulation, since the viscous resistance is proportional to U^2 , and hence decreases in effect at low speeds, the Coulomb type friction must become dominant at low speeds. However, in incorporating viscous intensification from material locking at low speeds, the importance of the μ type resistance should be less. In the case of avalanche flow the dominant effect of μ type friction should occur only at the start of the avalanche, when the motion is of a sliding type, until the snow blocks break up and a granular layer develops. From observations that avalanches seldom release on slopes less than 20 to 25°, a reasonable value of μ in the starting zone is $\mu = \tan 22^\circ = 0.4$.

In the terminal flow regime, we expect that with the locking mechanism operating, that μ should have values smaller than those in the range specified in the original Voellmy case. Also, from measurements by Lang, Dent (1982), frictional resistance is noted to be a linear function of the overburden load, so that the modified form of the friction coefficient, designated, μ^* , is

$$\mu^* = \mu_0 h$$

where μ_0 must be determined from numerical evaluation of avalanche flows.

In summary, the acceleration equation we now use for numerical evaluation of segment kinematics is

$$a = g(\sin \theta - \mu^* \cos \theta) - \frac{v^*}{R^3} v^2$$

where $\mu^* = \mu_0 h$ and $v^* = v_0(1 + C_0 e^{-1.25v})$. In regions where v^* is variable at low velocities, the segment length is decreased so that a step-wise linear approximation can be made of the acceleration, so that velocity at the end of each segment may still be computed using the equation

$$v_B = \left[\frac{g}{\beta} (1 - e^{-2\beta L}) + v_A^2 e^{-2\beta L} \right]^{1/2}$$

However, the computation of the final runout distance is no longer based upon evaluation of S_f , as in the Voellmy approach. Instead, with the grid refinement that is used, runout distance is simply the grid location where the velocity of flow becomes negligible.

PARAMETER EVALUATION: IRLINGTON PARK AVALANCHE

To proceed further, we evaluate C_0 , v_0 and μ_0 for a specific avalanche path for which the flow resistance is apparently constant over the entire path of runout. The Irlington Park path located in the San Juan Mountains in Colorado is used. Documentation of this path is given in detail by Lang, Dawson, Martinelli, 1979. For the Cheng-Perla evaluation (designated ACCEL hence forth) and the Biviscous-Equilibrium evaluation (designated BIEQ henceforth) an 11 segment approximation of the path was selected (Figure 9). The 11th segment is of a frozen lake bed, which is horizontal and extends for 350m, although only 50m of the segment is shown in Figure 9.

As a measure of comparison the ACCEL fit and the BIEQ fit are compared to corresponding results obtained from program AVALNCH. For example,

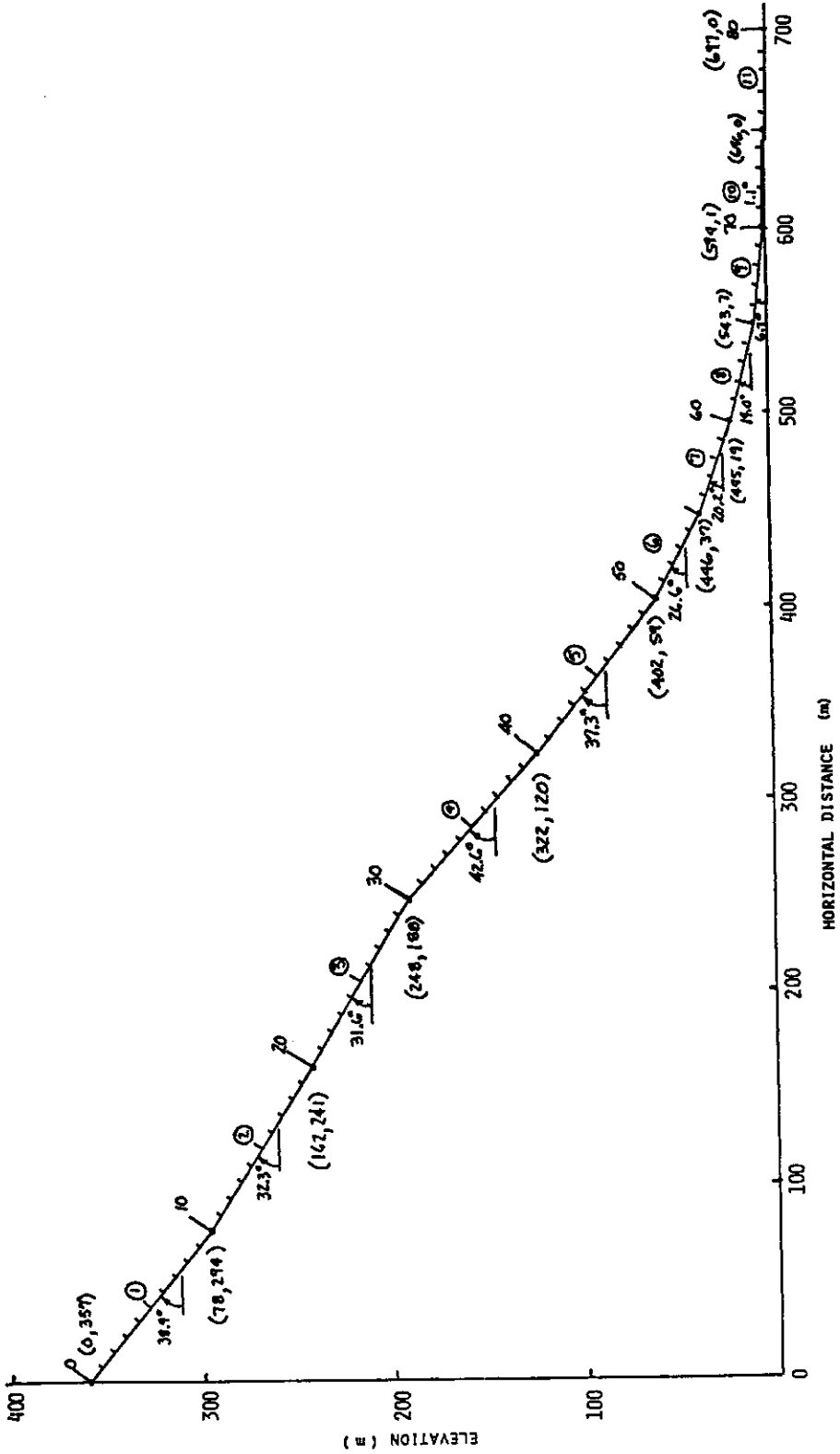


FIGURE 9: Ironton Park Avalanche Path

for a 2.0m nominal depth starting zone of snow, program AVALNCH computes a runout distance 230m into segment 11, and a maximum speed of approximately 42ms^{-1} . This same level of performance is obtained with the ACCEL program with $\mu=0.09$ and $M_D=400$. Note, that with program AVALNCH the released snow was distributed in cells 1 through 7 with a leading edge taper (Figure 10). In the ACCEL program computation the first segment was taken as the last 30m of segment ①.

To evaluate the BIEQ parameters, the Ironton Park avalanche path was evaluated first with program AVALNCH for starting zone snow depths of $h=0.5, 1.0, 1.5$ and 2.0m . Parameterization with program AVALNCH, was $f=0.5$ and $v=0.5\text{m}^2\text{s}^{-1}$. The four starting zone depth cases were then run with BIEQ, adjusting the governing parameters in order to obtain a best-fit of all the cases. The intent was to evaluate μ_o, v_o and C_o so that all the different depth cases were approximated by a single set of these parameters. With program AVALNCH the $h=0.5\text{m}$ case is a sluff onto the bench of cells 20 to 30 (Figure 9). In program BIEQ the same type of effect was modeled, but without considering a detailed duplication between the programs. A single set of parameters that models all cases was found to be $\mu_o=0.027, v_o=0.027$ and $C_o=500$. Using these values the maximum velocities and corresponding runout distances are summarized in Table 4. Program ACCEL was run only for the $h=2.0\text{m}$ case. Assuming that the primary function of these codes is to compute runout distance of the larger avalanches, then it is seen that all versions provide satisfactory results.

The distribution of velocity of the avalanche along the path can also be compared from the data that was obtained. The profiles are shown in Figures 11, 12 and 13 for $h=1.0, 1.5$ and 2.0m , respectively. It is seen that the BIEQ velocities are approximately 10 to 20 % higher than the AVALNCH values for the case $h=1.0\text{m}$, with runout distances

TABLE 4: Ironton Park Avalanche Path, computer program comparisons for different starting zone snow depths.

STARTING ZONE NOMINAL SNOW DEPTH (m)	MAXIMUM VELOCITY ALONG PATH (ms ⁻¹)			DISTANCE OF TRAVEL ALONG PATH (m)		
	AVALNCH	BIEQ	ACCEL	AVALNCH	BIEQ	ACCEL
	2.0	42.0	39.2	42.9	980	970
1.5	27.8	27.4	—	800	810	—
1.0	13.0	15.4	—	710	720	—
0.5	—	—	—	SLUFF	SLUFF	—

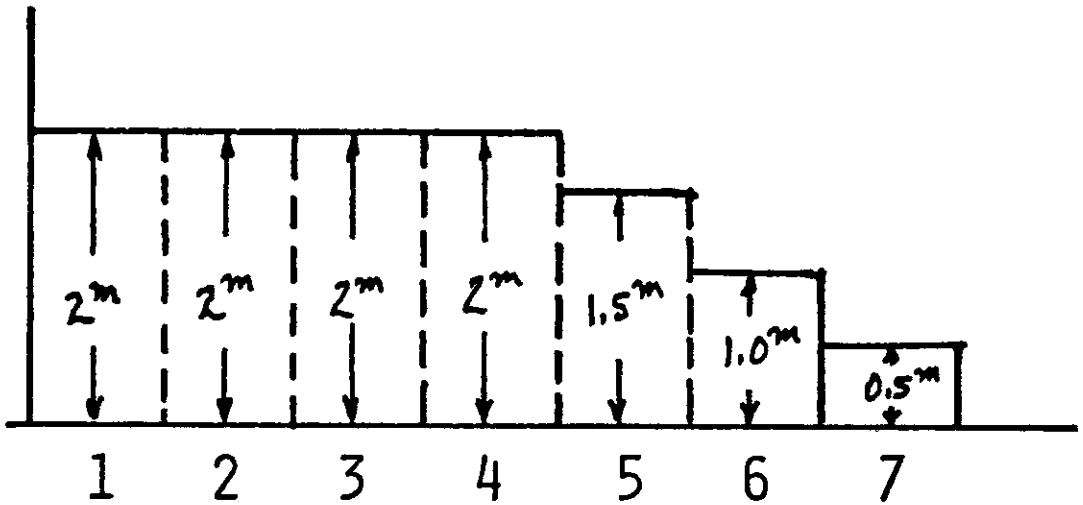


FIGURE 10: SNOW DISTRIBUTION IN THE STARTING ZONE OF THE IRONTON PARK AVALANCHE PATH.

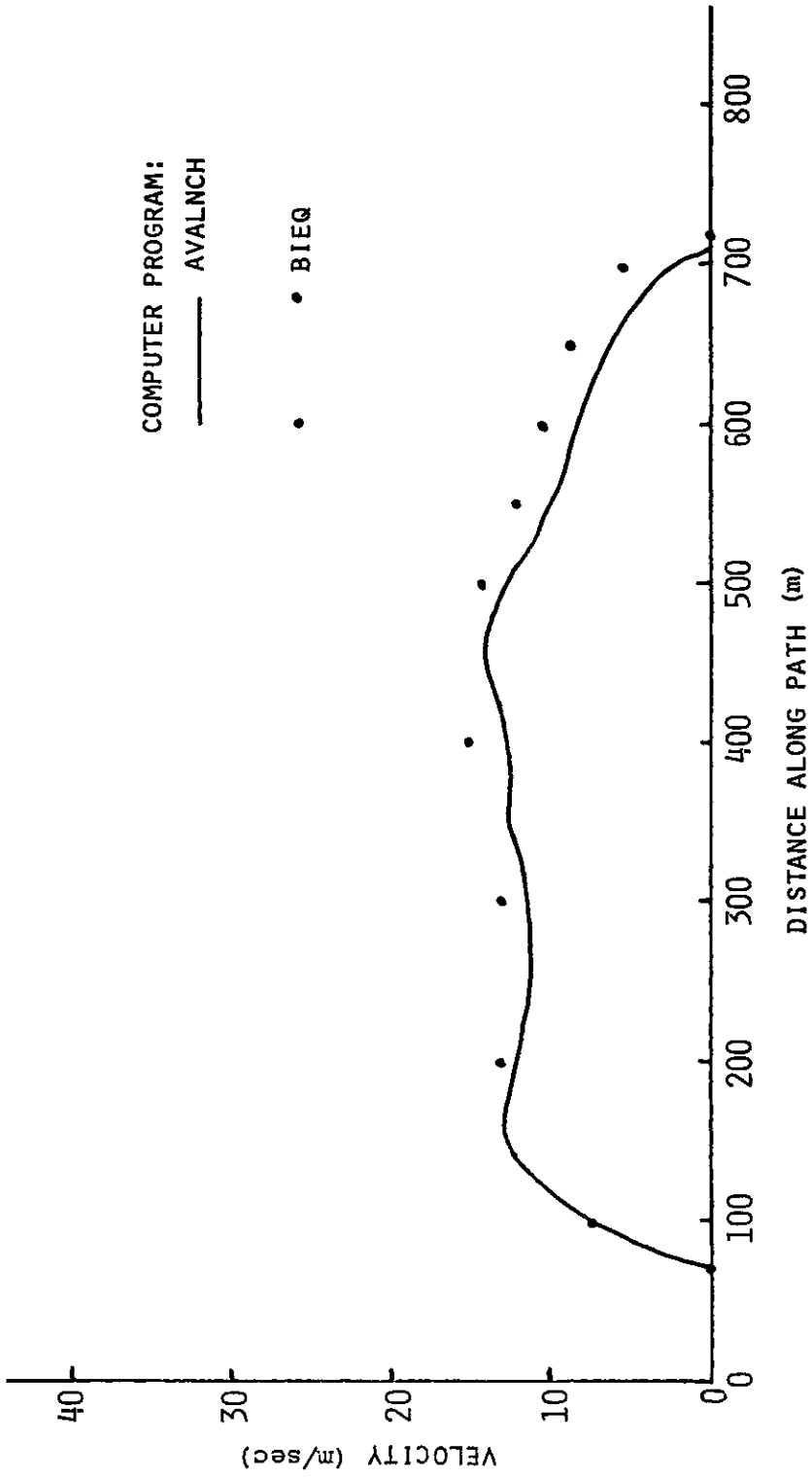


FIGURE 11: IRONTON PARK AVALANCHE PATH : VELOCITY PROFILE FOR STARTING ZONE SNOW DEPTH $h_s = 1.0m$.

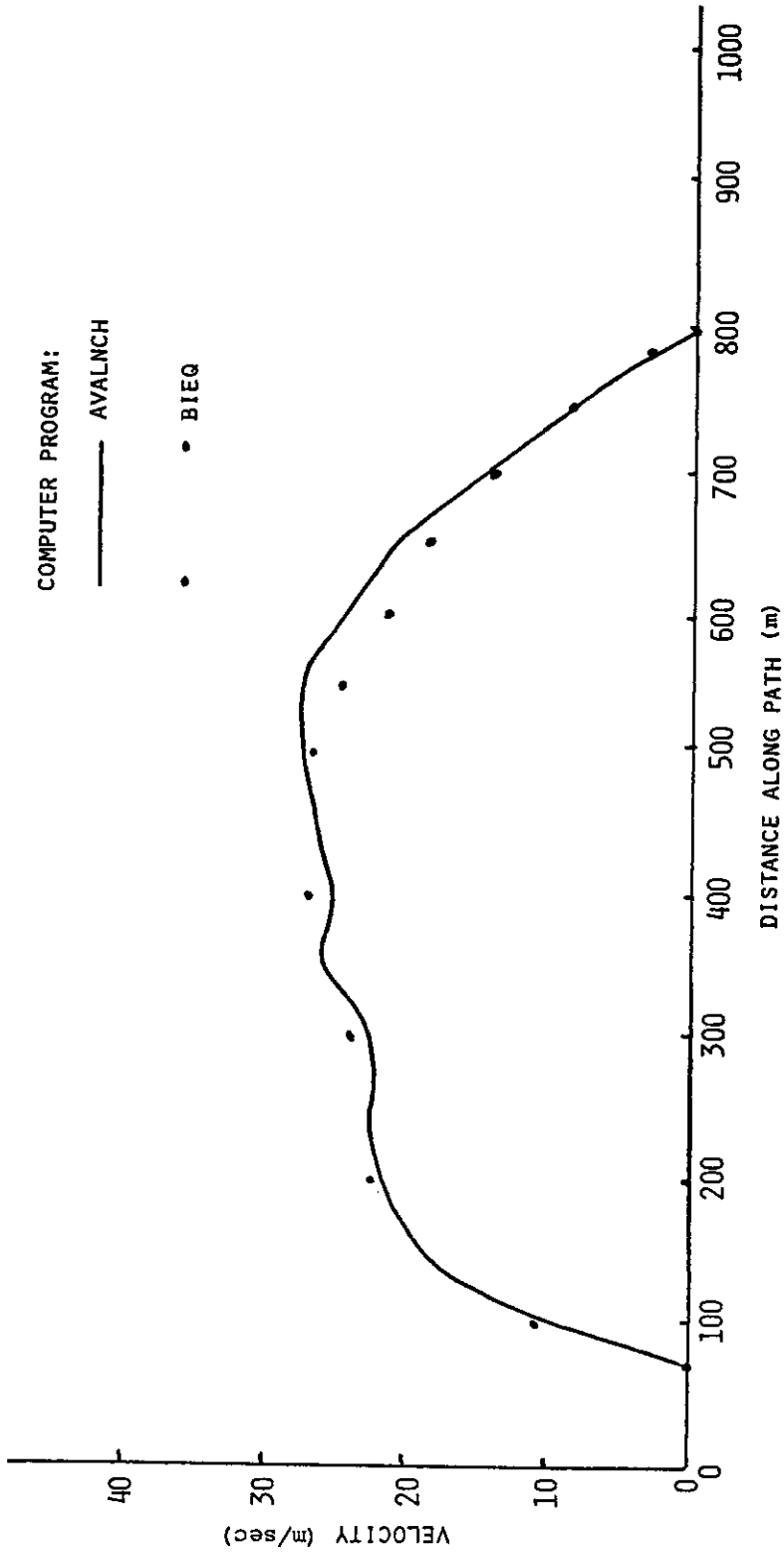


FIGURE 12: IRONTON PARK AVALANCHE PATH : VELOCITY PROFILE FOR STARTING ZONE SNOW DEPTH $\bar{h} = 1.5m$.

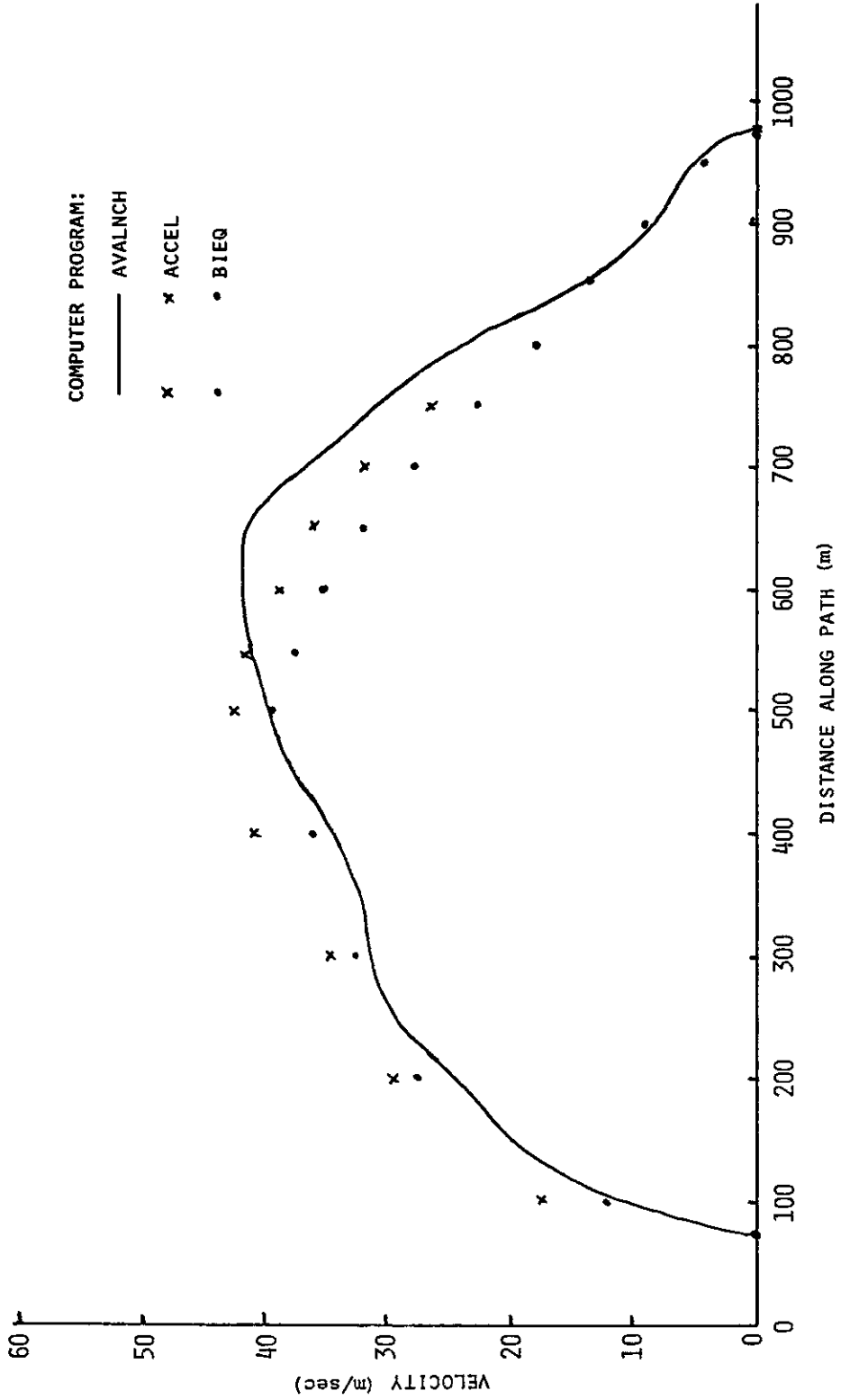


FIGURE 13: IRONTON PARK AVALANCHE PATH : VELOCITY PROFILE FOR STARTING ZONE SNOW DEPTH $h_s = 2.0m$.

approximately the same. The closest fit between the two sets of data is obtained for the $h = 1.5\text{m}$ case, where differences are less than 10%. For the case of a starting zone snow depth of $h = 2.0\text{m}$, the shapes of the profiles are different, with attendant larger differences at some data points (up to 30%). Shown also for this case is a profile from program ACCEL. If modeling of the other cases is carried out with program ACCEL, different values of parameters M/ρ and μ would be necessary, as no explicit dependence on flow depth is included in the parameterization of this program. Parameterization in all of these computer results is not unique, and it is likely that different shapes would be obtained if different, but equally valid, parameterization is used. In the case of program AVALNCH the parameterization used has been correlated with an actual experimental velocity profile, which, however, was not of the Ironton Park path (LaChapelle, Lang, 1980). Since program AVALNCH is based upon transient fluid motion, while BIEQ (and ACCEL) is based upon equilibrium fluid dynamics, it should not be expected that the profiles be in complete agreement. The numerical experimentation necessary in order to determine if other parameterization of BIEQ (or of ACCEL) produce better profile correspondance does not seem warranted at the present time, since detailed experimental data is first needed in order to establish an absolute basis of comparison.

With ν_0 and C_0 now known for the Ironton Park avalanche path, the material locking viscous equation is

$$\nu^* = \nu_0 \left(1 + 500 e^{-1.25 U} \right)$$

where $\nu_0 = 0.027\text{m}^2$. From this equation, viscosity ν^* varies with velocity, as shown in Figure 14. Viscous drag, proportional to $\nu^* U^2$, increases from $U = 5\text{ms}^{-1}$ to about $U = 2\text{ms}^{-1}$, whereas the ordinary drag $\nu_0 U^2$ monotonically decreases with considerably smaller magnitude. From $U = 2\text{ms}^{-1}$ to $U = 1\text{ms}^{-1}$ the viscous drag $\nu^* U^2$ begins to decrease, as the U^2 factor

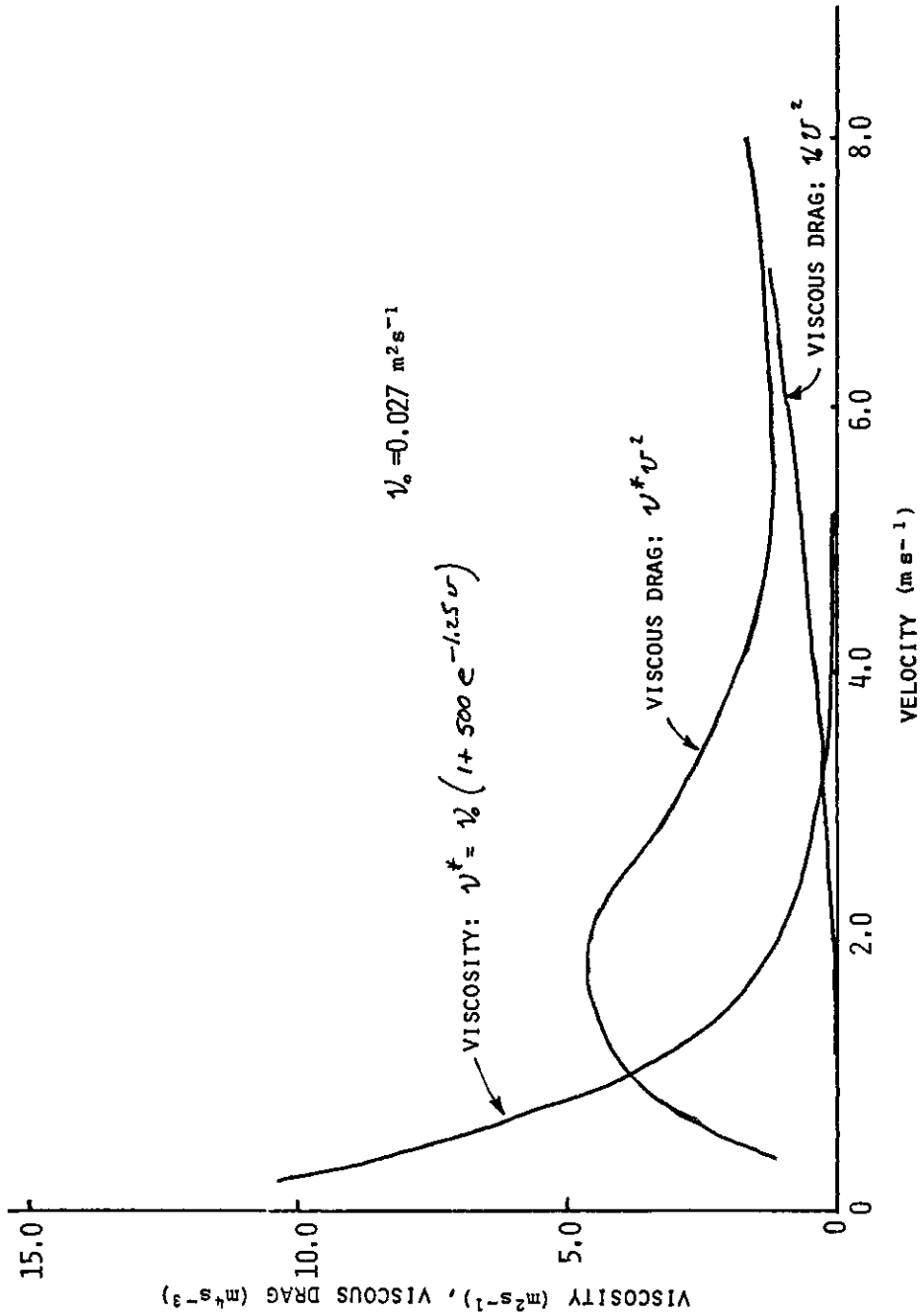


FIGURE 14: VISCOSITY AND VISCIOUS DRAG VERSE FLOW VELOCITY

dominates over the exponential form of v^* . However, for the larger avalanches for which these computer programs are assumed applicable, $v < 1 \text{ ms}^{-1}$ is a small velocity of negligible order. Thus, in program BIEQ if the velocities v_A and v_B at the beginning and end of a segment are both less than 1.0 ms^{-1} , then computations are terminated.

FLOW PARAMETER SENSITIVITIES

The basic parameters of program BIEQ are assumed to be the viscosity ν_0 , friction coefficient μ_0 , and flow depth h . We look now at the sensitivity on runout distance to small variations in these parameters. Likewise, the parameters of program ACCEL are M/D and μ , for which small variations are considered with respect to runout. In these calculations the Ironton Park avalanche path is used, and the reference avalanche configuration is taken as $h = 2.0 \text{ m}$, with a runout distance into segment (ii) of 230m. Results are presented for small percentage variations in the basic parameters (Figure 15). With program BIEQ the depth of the snow release is the most sensitive, by a factor of 2, of the parameters considered. Next in order of sensitivity is viscosity, followed by friction at about half that of viscosity. Starting zone snow depth is not explicit in program ACCEL so no evaluation is given. However, for viscosity and friction the sensitivities are in general correspondence with those of BIEQ.

Although the Ironton Park avalanche path has no negative or adverse slope along its' length, to check sensitivity of this parameter, the flat runout segment (ii) was given representative negative slope values. At a slope angle of -10° runout was 40m into segment (ii). For small percentage changes in this angle the percent change in runout is of the same order as that for viscosity (Figure 15).

The results of this sensitivity study show that flow depth is a

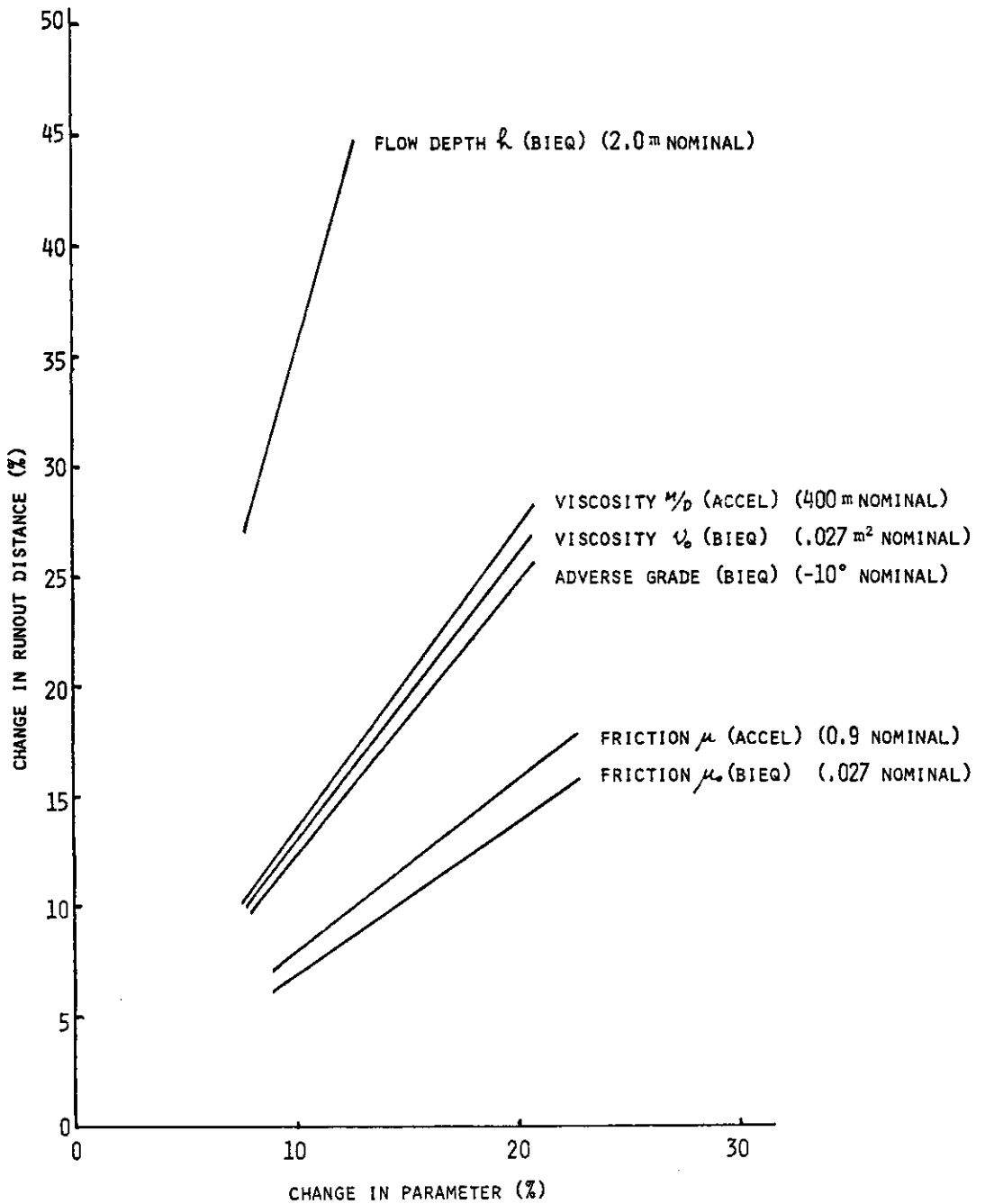


FIGURE 15: CHANGE IN AVALANCHE RUNOUT DISTANCE WITH CHANGE IN VARIOUS FLOW PARAMETERS FOR THE IRONTON PARK AVALANCHE.

variable of primary importance in avalanche runout prediction. With no explicit representation of h in the form of the Voellmy equation used in ACCEL, it requires large variation of the parameters in this equation in order to model the varying conditions of different avalanches and different avalanche paths. With the explicit representation of h in program BIEQ, the variation of the remaining coefficients in order to model different avalanche paths should be less. No claim is intended that in program BIEQ that the h dependence is exact. However, because of the inherent sensitivity of h in avalanche runout, it should be considered, and experimentation should be carried out to define more rational dependence of h in the avalanche equations of motion.

COMPUTER PROGRAM BIEQ

The version of computer program BIEQ listed in Table 5 has array dimensions that allow up to 100 slope segments to be input. The order and format of the input data is summarized as follows:

Line 1: FORMAT (I10)

Columns 1-10: IC - Integer number of test cases that are to be run (right adjusted)

Line 2: FORMAT (4QA2)

Columns 1-80: Name and identification information for test case #1

Line 3: FORMAT (4F10.0)

Columns 1-10: H - Snow depth in the starting zone (m).

Columns 11-20: YNU - high stress viscosity (m^2). If the value is set to zero here, it indicates to the computer that an array of values will be input (Line 4).

Columns 21-30: SML - sub-element length (m).

Columns 31-40: YMU - friction coefficient

Line 4: FORMAT (I10, 3F10.0)

Columns 1-10: IS - segment number
Columns 11-20: ANGLE - segment slope (°).
Columns 21-30: SEGL - segment length (m).
Columns 31-40: VNU - segment viscosity (m²).
If YNU=0.0 in line 3, then
viscosity must be input for
each segment. If YNU > 0.0
then any values input here
are disregarded.

The format of line 4 is repeated for as many segments as are used to represent an avalanche path. After all segments have been listed a blank card should follow. Following the blank card a second set of data may be input in the same order and format as Lines 2 through 4, until a number of cases equal to the value of IC of Line 1 have been set up. If the program is to be used to run only one case at a time, then statements 5, 6, 21, 105 and 106 may be eliminated from the program (Table 5).

Regarding other parameters in the program, the value of friction in the first segment of flow of an avalanche is defined in statement 49 by YUU=0.4, which may be changed at user discretion. Statement 61 is the test for negligible flow speed, that if $U_A < 1.0\text{ms}^{-1}$ and $U_B < 1.0\text{ms}^{-1}$, the computations are terminated. Error messages following statement 92 account for the following conditions:

- 1) The avalanche does not stop, and calculations are terminated.
- 2) The number of segments exceeds IMAX=100, and calculations are terminated.
- 3) The flow velocity through a segment is negligible, and calculations are terminated.
- 4) Viscosity is not specified, by one of two possible input options, and calculations are terminated.

The segment mini-length parameter SML, input to the program by statement 11, was taken as 10.0m for the Ironton Park path analysis, which

provided sufficient accuracy in runout distance prediction. For different avalanche paths, user option is to change the value of SML. When the program switches to the mini-segment analysis option, the velocity at the end of each mini-segment increment is printed to the right of the mainline printout. Thus, the user knows when the switch has been made in the program. The mainline output from the program consists of a listing of segment number, velocity V_A at the start of the segment, and velocity V_B at the end of the segment. If the avalanche stops within a segment, then the output is the segment number in which the avalanche stops, the velocity, V_A , at the start of the segment, and the total runout distance measured along the path, which is the sum of all segment lengths up to the stopping point. The partial segment length of the segment in which the avalanche stops is also included in the sum. Also output by the program, for reference purposes, are the values of all input parameters.

TABLE 5: Listing of Computer Program BIEQ.

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```

1      C      PROGRAM BIEQ
2      DIMENSION NAME(40),SEGL(100),THETA(100),VNU(100)
3      IMAX=100
4      G=9.806
5      READ(7,9) IC
6      IF=1
7      5 READ(7,20) NAME
8      WRITE(6,30) NAME
9      WRITE(6,40)
10     C      READ INPUT DATA
11     READ(7,50) H,YNU,SML,YMU
12     I=1
13     70 READ(7,10) IS,ANGLE,SEGL(I),VNU(I)
14     IF(YNU.GT.0.0001) VNU(I)=YNU
15     IF(IS.EQ.0) GO TO 100
16     WRITE(6,80) IS,ANGLE,SEGL(I),VNU(I)
17     THETA(I)=3.14159*ANGLE/180.0
18     I=I+1
19     IF(I.GT.IMAX) GO TO 444
20     GO TO 70
21     9 FORMAT(I10)
22     10 FORMAT(I10,3F10.0)
23     20 FORMAT(40A2)
24     30 FORMAT(1H1,5X,40A2)
25     40 FORMAT(1H0,15X,'INPUT DATA'//10X,'SEGMT ANGLE SEGL VSCSTY')
26     50 FORMAT(4F10.0)
27     60 FORMAT(1H0,9X,'SNOW DEPTH=',F5.2/10X,'HIGH STRESS VISCSTY=',
28     *F6.4/10X,'SEGMENT MINI-LNGTH=',F6.2/10X,'FRICTION COEF=',F5.3)
29     80 FORMAT(10X,I4,2X,F8.1,1X,F6.1,2X,F6.4)
30     210 FORMAT(1H0,20X,'RESULTS'//10X,'SEGMT',5X,'VA',7X,'VB')
31     240 FORMAT(5X,I9,F10.2,F9.2)
32     260 FORMAT(5X,I9,F10.2,3X,'RUNOUT=',F10.2)
33     270 FORMAT(50X,'V=',F9.3)
34     334 FORMAT(10X,'AVALANCHE DOES NOT STOP')
35     445 FORMAT(5X,'SEGMENT NUMBER EXCEEDS SPECIFIED IMAX')
36     556 FORMAT(10X,'FLOW VELOCITY NEGLIGIBLE')
37     667 FORMAT(10X,'VISCOSITY NOT SPECIFIED')
38     888 FORMAT(10X,'COMPUTATIONS TERMINATED')
39     C
40     C      INITIAL COMPUTATIONS
41     100 WRITE(6,60) H,YNU,SML,YMU
42     IF(VNU(1).LE.0.0) GO TO 666
43     VA=0.0
44     IT=I-1
45     I=1
46     120 IM=1
47     SUL=SML
48     YUU=YMU
49     IF(I.EQ.1) YUU=0.4
50     IF(I.EQ.1) WRITE(6,210)
51     A=G*SIN(THETA(I))-G*YUU*H*COS(THETA(I))
52     B=VNU(I)*(1.0+500.*EXP(-1.25*VA))/H**3
53     IF(I.EQ.1) A=G*SIN(THETA(I))-G*YUU*COS(THETA(I))
54     IF(I.EQ.1) B=2.0*VNU(I)/H**3

```

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```

55      E=EXP(-2.0*B*SEGL(I))
56      F=VA*VA*E+A*(1.0-E)/B
57      IF(P.LT.64.0) GO TO 150
58      130 VB=SQRT(P)
59      WRITE(6,240) I,VA,VB
60      Q=THETA(I)-THETA(I+1)
61      IF(VA.LT.1.0.AND.VB.LT.1.0) GO TO 555
62      VA=VB*COS(Q)
63      IF(Q.LE.0.0) VA=VB
64      I=I+1
65      IF(I.GT.IT) GO TO 333
66      GO TO 120
67      C    SUB-SEGMENT COMPUTATIONS
68      150 IN=INT(SEGL(I)/SUL)
69      V=VA
70      SL=SEGL(I)
71      SEGL(I)=0.0
72      IF(IN.LE.1) GO TO 180
73      170 B=VNU(I)*(1.0+500.*EXP(-1.25*V))/H**3
74      IF(VA.EQ.0.0) B=2.0*VNU(I)/H**3
75      E=EXP(-2.0*B*SUL)
76      F=V*V*E+A*(1.0-E)/B
77      SEGL(I)=SEGL(I)+SUL
78      IF(P.LE.0.0) GO TO 190
79      IF(IM.EQ.IN+1) GO TO 130
80      V=SQRT(P)
81      WRITE(6,270) V
82      IF(IM.EQ.IN) SUL=SL-SEGL(I)
83      IM=IM+1
84      GO TO 170
85      180 SEGL(I)=SL
86      IF(P) 190,190,130
87      190 S=0.0
88      DO 200 J=1,I
89      200 S=S+SEGL(J)
90      WRITE(6,260) I,VA,S
91      GO TO 999
92      C    ERROR MESSAGES
93      333 WRITE(6,334)
94      WRITE(6,888)
95      GO TO 999
96      444 WRITE(6,445)
97      WRITE(6,888)
98      GO TO 999
99      555 WRITE(6,556)
100     WRITE(6,888)
101     GO TO 999
102     666 WRITE(6,667)
103     WRITE(6,888)
104     999 CONTINUE
105     IP=IP+1
106     IF(IP.NE.IC+1) GO TO 5
107     STOP
108     END

```

Computer Program ACEL

Computer program ACCEL, developed by Cheng and Perla (1979), is listed with detailed explanations in their reference. A modification of their program, designated ACEL, that operates on the Melcom 70 Computer, is listed in Table 6. Since the ACCEL and ACEL programs are similar, no descriptive summary is included herein for program ACEL.

TABLE 6: Listing of Computer Program ACEL.

```

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1 C PROGRAM ACEL
2 DIMENSION THETA(100), SLENG(100), NAME(40), YMU(100), YMD(100)
3 READ(7,80) IC
4 IMAX=100
5 G=9.806
6 IP=0
7 5 WRITE(6,10)
8 10 FORMAT(1H1)
9 READ(7,20) NAME
10 20 FORMAT(40A2)
11 WRITE(6,30) NAME
12 30 FORMAT(5X,40A2)
13 READ(7,40) FMU,FMD
14 40 FORMAT(F10.0,F10.0)
15 IF(FMU.GT.0.0.OR.FMD.GT.0.0) WRITE(6,50) FMU,FMD
16 50 FORMAT(13X,'MU=',F5.2,5X,'M/D=',F8.0)
17 WRITE(6,60)
18 60 FORMAT(1H0,15X,'INPUT DATA')
19 WRITE(6,70)
20 70 FORMAT(1H0,10X,'SEGMENT ANGLE LENG MU M/D')
21 210 FORMAT(1H0,20X,'RESULTS'/10X,'SEGMENT',5X,'VA',7X,'VB')
22 240 FORMAT(5X,I9,F10.2,F9.2)
23 300 FORMAT(5X,I9,F10.2,3X,'RUNOUT=',F10.2)
24 I=1
25 75 READ(7,80) IS,ANGLE,SLENG(I),YMU(I),YMD(I)
26 IF(IS.EQ.0) GO TO 100
27 80 FORMAT(I10,4F10.0)
28 IF(YMU(I).LE.0.0) YMU(I)=FMU
29 IF(YMD(I).LE.0.0) YMD(I)=FMD
30 IF(YMD(I).EQ.0.0) GO TO 666
31 WRITE(6,90) IS,ANGLE,SLENG(I),YMU(I),YMD(I)
32 90 FORMAT(10X,I4,2X,F8.1,1X,F6.1,F6.2,1X,F6.0)
33 THETA(I)=ANGLE*3.1416/180.
34 I=I+1
35 IF(I.GT.IMAX) GO TO 777
36 GO TO 75
37 100 II=I-1
38 I=1
39 215 VA=0.0
40 220 ALPHA=G*SIN(THETA(I))-G*YMU(I)*COS(THETA(I))
41 PA=EXP(-2.0*SLENG(I)/YMD(I))
42 P=VA*VA*PA+ALPHA*YMD(I)*(1.0-PA)
43 IF(P) 290,280,230
44 230 VB=SQRT(P)
45 IF(I.EQ.1) WRITE(6,210)
46 WRITE(6,240) I,VA,VB
47 Q=THETA(I)-THETA(I+1)
48 IF(Q) 260,260,250
49 250 VA=VB*COS(Q)
50 GO TO 270
51 260 VA=VB
52 270 I=I+1
53 IF(I.GT.II) GO TO 333
54 GO TO 220

```


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```
55      280 VB=0.0
56      WRITE(6,240) I,VA,VB
57      I=I+1
58      IF(I.GT.II) GO TO 333
59      GO TO 215
60      290 IF(ALPHA.EQ.0.0) GO TO 555
61      DD=1.0-(VA*VA)/(ALPHA*YMD(I))
62      IF(DD.LT.0.0) GO TO 444
63      S=0.5*YMD(I)*ALOG(DD)
64      WRITE(6,300) I,VA,S
65      GO TO 999
66      333 WRITE(6,334)
67      334 FORMAT(10X,'AVALANCHE DOES NOT STOP')
68      WRITE(6,888)
69      GO TO 999
70      444 WRITE(6,445)
71      445 FORMAT(5X,'ARGUMENT OF LOG CANNOT BE NEGATIVE')
72      WRITE(6,888)
73      GO TO 999
74      555 WRITE(6,556) I
75      556 FORMAT(5X,'ALPHA IS ZERO FOR SEGMENT',I4)
76      WRITE(6,888)
77      GO TO 999
78      666 WRITE(6,667) I
79      667 FORMAT(5X,'M/D IS ZERO FOR SEGMENT',I4)
80      WRITE(6,888)
81      GO TO 999
82      777 WRITE(6,778)
83      778 FORMAT(5X,'SEGMENT NUMBER EXCEEDS SPECIFIED IMAX')
84      WRITE(6,888)
85      888 FORMAT(10X,'COMPUTATIONS TERMINATED')
86      999 IP=IP+1
87      IF(IP.NE.IC) GO TO 5
88      STOP
89      END
```

PROGRAM BIEQ option studies

Computer program BIEQ was changed in several ways in order to evaluate other possible options of the code. One change was to eliminate dry friction and to incorporate a true biviscous material representation. The ratio between the low stress and high stress viscosities was set at 10, and the transition velocity between the two viscosities was set at $U_T = 5$ and 8 ms^{-1} in different evaluations. Results are summarized in Table 7. While runout distances can be duplicated to errors less than 5% by appropriate selection of the high stress viscosity, the predicated maximum runout speeds in all cases are low by as much as 25%. This indicates that viscosity, when set for correct runout, is too large for sufficient runout speed. Increasing the low stress range from $U_T = 5$ to 8 ms^{-1} reduced the maximum speed error to 18% or less. However, the largest error in max speed is with the 2.0m deep avalanche, and it is with the deeper avalanches that accurate modeling is wanted. These results were obtained with a low speed cutoff of $U = 1.0 \text{ ms}^{-1}$ for computer termination of calculations. In other program versions the low speed cutoff is expressed as a percentage of maximum speed of the avalanche. The next modification of BIEQ was to incorporate a percent cutoff, the value set at 7.0%. Thus, when the speed of each avalanche dropped below 7% of maximum speed over an entire segment, computations were terminated. Results based upon this program option are summarized in Table 8. It is noted that runout distances are better matched with the transition speed $U_T = 5.0 \text{ ms}^{-1}$ compared with $U_T = 8.0 \text{ ms}^{-1}$. Putting a percent low speed cutoff reduced the error in max. speed from 25% to 22%, but remains a large error. Also evaluated was increasing the viscosity ratio from 10 to 20; however the effect of this was insignificant in changing any of the kinematical data.

The results of these option studies with program BIEQ are not yielding

TABLE 7: Ironton Park avalanche study using program BIEQ with no dry friction and a true biviscous material representation.

Depth of flow (m)	trans. sp: $U_T = 5.0\text{ms}^{-1}$ viscosity: $\nu = .05\text{m}^2$		trans. sp: $U_T = 8.0\text{ms}^{-1}$ viscosity: $\nu = .04\text{m}^2$		PROGRAM AVALNCH $\nu = 0.23\text{m}^2$	
	max vel. (ms ⁻¹)	runout distance (m)	max vel. (ms ⁻¹)	runout distance (m)	max vel. (ms ⁻¹)	runout distance (m)
1.0	11.5	650	12.9	650	12.0	640
1.5	21.0	720	23.4	700	28.0	730
2.0	31.3	900	34.4	910	42.0	910

TABLE 8: Ironton Park avalanche study using program BIEQ with no dry friction, a true biviscous material representation, and a 7% low speed cutoff option.

Depth of flow (m)	trans. sp: $U_T = 5.0\text{ms}^{-1}$ viscosity: $\nu = 0.045\text{m}^2$		trans. sp: $U_T = 8.0\text{ms}^{-1}$ viscosity: $\nu = 0.037\text{m}^2$		PROGRAM AVALNCH $\nu = 0.23\text{m}^2$	
	max vel. (ms ⁻¹)	runout distance (m)	max vel. (ms ⁻¹)	runout distance (m)	max vel. (ms ⁻¹)	runout distance (m)
1.0	12.1	640	13.4	650	12.0	640
1.5	22.1	720	24.3	700	28.0	730
2.0	32.8	920	35.6	910	42.0	910

as satisfactory a fit to avalanche runout as was obtained with the version of BIEQ with small dry friction and a gradual biviscous transition at low speeds. An abrupt change in viscosity at a transition speed is unlikely to be physically accurate, and with only one segment modeling in BIEQ, velocity changes rapidly when the viscosity changes. However, the gradual viscosity change used previously in BIEQ, also is not physically based, and may not be generally applicable to different avalanche problems.

A listing of program BIEQ with no dry friction, true biviscous material representation, and 7% velocity cutoff options is given in Table 9. The program was given the code name BEAR with these option changes.

Another modification incorporates the improved results obtained from previous studies, and uses a true biviscous model coupled with low friction. The previous improved results were with a transition speed $U_T = 5.0 \text{ms}^{-1}$, and a velocity cutoff at 5% of maximum velocity. Findings of this evaluation are summarized in Table 10.

TABLE 10: Ironton Park avalanche study using program BIEQ with a true biviscous material representation, low dry friction, and a 5% low speed cutoff option.

Depth of flow (m)	trans. speed : $U_T = 5.0 \text{ms}^{-1}$ viscosity : $\nu = 0.025 \text{m}^2$ friction : $\mu = 0.030$		Program AVALNCH $\nu = 0.23 \text{m}^2$	
	maximum velocity (ms^{-1})	runout distance (m)	maximum velocity (ms^{-1})	runout distance (m)
1.0	16.0	660	14.0	640
1.5	28.3	740	28.0	730
2.0	40.4	910	41.0	910

These results are comparable to the results obtained using a gradually changing viscosity through the transition speed range. The inclusion of small dry friction coupled with viscosity allows matching of both runout distance and max velocity for the different release depth cases.

TABLE 9: Listing of computer program BIEQ under option changes
of (a) no dry friction, (b) true biviscous material
representation, and (c) 7% velocity cutoff.

INE.	1	2	3	4	5	6	7
1	C	PROGRAM BEAR: VARIABLE TRANSITION VEL.: VEL. THRESHOLD 7%					
2		DIMENSION NAME(40),SEGL(100),THETA(100),VNU(100)					
3		IMAX=100					
4		G=9.806					
5		5	READ(7,20) NAME				
6			WRITE(6,30) NAME				
7			WRITE(6,40)				
8	C	READ INPUT DATA					
9		READ(7,50) H, YNU, SML, XNU, VP					
10		I=1					
11		70	READ(7,10) IS, ANGLE, SEGL(I), VNU(I)				
12		IF (YNU.GT.0.0) VNU(I)=YNU					
13		IF (IS.EQ.0) GO TO 100					
14		WRITE(6,80) IS, ANGLE, SEGL(I), VNU(I)					
15		THETA(I)=3.14159*ANGLE/180.0					
16		I=I+1					
17		IF (I.GT.IMAX) GO TO 444					
18		GO TO 70					
19		9	FORMAT(I10)				
20		10	FORMAT(I10,3F10.0)				
21		20	FORMAT(40A2)				
22		30	FORMAT(1H1,5X,40A2)				
23		40	FORMAT(1H0,15X,'INPUT DATA'//10X,'SEGMT ANGLE SEGL VSCSTY')				
24		50	FORMAT(5F10.0)				
25		60	FORMAT(1H0,9X,'SNOW DEPTH=',F5.2/10X,'SEGMT MINI-LNSTH=',F6.2,				
26			*/10X,'VSCSITY MULT FACTOR=',F5.1/10X,'VEL AT TRANSITION=',F5.1)				
27		80	FORMAT(10X,I4,2X,F8.1,1X,F6.1,2X,F6.4)				
28		210	FORMAT(1H0,20X,'RESULTS'//10X,'SEGMT',5X,'VA',7X,'VB',				
29			*9X,'S',9X,'T')				
30		240	FORMAT(5X,I9,F10.2,F9.2,F9.2,F9.2)				
31		260	FORMAT(5X,I9,F10.2,3X,'RUNOUT=',F10.2)				
32		270	FORMAT(53X,'V=',F9.3)				
33		334	FORMAT(10X,'AVALANCHE DOES NOT STOP')				
34		445	FORMAT(5X,'SEGMENT NUMBER EXCEEDS SPECIFIED IMAX')				
35		556	FORMAT(10X,'FLOW VELOCITY NEGLIGIBLE')				
36		667	FORMAT(10X,'VISCOSITY NOT SPECIFIED')				
37		888	FORMAT(10X,'COMPUTATIONS TERMINATED')				
38	C						
39	C	INITIAL COMPUTATIONS					
40		100	WRITE(6,60) H,SML,XNU,VP				
41		IF (VNU(1).LE.0.0) GO TO 666					
42		VA=0.0					
43		VMAX=0.0					
44		IT=I-1					
45		I=1					
46		S=0.0					
47		T=0.0					
48		120	IM=1				
49		SUL=SML					
50		IF (I.EQ.1) WRITE(6,210)					

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LINE.	1	2	3	4	5	6	7
	12345678901234567890123456789012345678901234567890123456789012345678901						
51		A=G*SIN(THETA(I))					
52		B=VNU(I)/H**3					
53		E=EXP(-2.0*B*SEGL(I))					
54		F=VA*VA*E+A*(1.0-E)/B					
55		VP2=VP*VP					
56		IF(VA.ER.0.0.AND.F.LT.VP2) GO TO 130					
57		IF(VA.NE.0.0.AND.F.LT.VP2) GO TO 150					
58	130	VB=SQRT(F)					
59		IF(VMAX.LT.VB) VMAX=VB					
60		S=S+SEGL(I)					
61		T=T+2.0*SEGL(I)/(VA+VB)					
62		WRITE(6,240) I,VA,VB,S,T					
63		Q=THETA(I)-THETA(I+1)					
64		VT=0.07*VMAX					
65		IF(VA.LT.VT.AND.VB.LT.VT) GO TO 190					
66		VA=VB*COS(Q)					
67		IF(Q.LE.0.0) VA=VB					
68		I=I+1					
69		IF(I.GT.IT) GO TO 333					
70		GO TO 120					
71	C	SUB-SEGMENT COMPUTATIONS					
72	150	IN=INT(SEGL(I)/SUL)					
73		V=VA					
74		SL=SEGL(I)					
75		SEGL(I)=0.0					
76		IF(IN.LE.1) GO TO 180					
77	170	IF(V.LE.VP) B=XNU*VNU(I)/H**3					
78		IF(V.GT.VP) B=VNU(I)/H**3					
79		E=EXP(-2.0*B*SUL)					
80		P=V*V*E+A*(1.0-E)/B					
81		VN=SQRT(P)					
82		IF(VMAX.LT.VN) VMAX=VN					
83		VT=0.07*VMAX					
84		SEGL(I)=SEGL(I)+SUL					
85		IF(VN.LT.VT.AND.V.LT.VT) GO TO 190					
86		IF(IM.EQ.IN+1) GO TO 130					
87		V=VN					
88		WRITE(6,270) V					
89		IF(IM.EQ.IN) SUL=SL-SEGL(I)					
90		IM=IM+1					
91		GO TO 170					
92	180	SEGL(I)=SL					
93		GO TO 130					
94	190	S=0.0					
95		DO 200 J=1,I					
96	200	S=S+SEGL(J)					
97		WRITE(6,260) I,VA,S					
98		GO TO 555					
99	C	ERROR MESSAGES					
100	333	WRITE(6,334)					

LINE.	1	2	3	4	5	6	7
	1234567890123456789012345678901234567890123456789012345678901						
101							WRITE (6,888)
102							GO TO 999
103	444						WRITE (6,445)
104							WRITE (6,888)
105							GO TO 999
106	555						WRITE (6,556)
107							WRITE (6,888)
108							GO TO 999
109	666						WRITE (6,667)
110							WRITE (6,888)
111	999						CONTINUE
112							STOP
113							END

*** PRINT END ***

For the three cases, maximum error in velocity is 14%, and maximum error in runout is 3%, both for the $h=1.0m$ case. For the larger release depths the errors reduce, which is a desirable result, as the intention is to better model the deeper flows.

Table 11 is a listing of the computer program that incorporates the improvements described above.

TABLE 11: Listing of program BIEQ with options of (a) low dry friction, (b) true biviscous material representation, and (c) 5% velocity cutoff.

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1      C PROGRAM BIEQ WITH TRUE BIVISCOUS, 5% CUTOFF VELOCITY OPTIONS
2      DIMENSION NAME(40),SEGL(100),THETA(100),VNU(100)
3      IMAX=100
4      G=9.806
5      5 READ(7,20) NAME
6      WRITE(6,30) NAME
7      WRITE(6,40)
8      C READ INPUT DATA
9      READ(7,50) H,YNU,SML,XNU,YMU
10     I=1
11     70 READ(7,10) IS,ANGLE,SEGL(I),VNU(I)
12     IF(YNU.GT.0.0) VNU(I)=YNU
13     IF(IS.EQ.0) GO TO 100
14     WRITE(6,80) IS,ANGLE,SEGL(I),VNU(I)
15     THETA(I)=3.14159*ANGLE/180.0
16     I=I+1
17     IF(I.GT.IMAX) GO TO 444
18     GO TO 70
19     10 FORMAT(1I0,3F10.0)
20     20 FORMAT(40A2)
21     30 FORMAT(1H1,5X,40A2)
22     40 FORMAT(1H0,15X,'INPUT DATA'//10X,'SEGMT ANGLE SEGL VSCSTY')
23     50 FORMAT(5F10.0)
24     60 FORMAT(1H0,9X,'SNOW DEPTH=',F5.2/10X,'SEGMT MINI-LNGTH=',F6.2,
25     */10X,'VSCSITY MULT FACTOR=',F5.1/10X,'FRICTION COEF=',F5.3)
26     80 FORMAT(10X,14,2X,F8.1,1X,F6.1,2X,F6.4)
27     210 FORMAT(1H0,20X,'RESULTS'/10X,'SEGMT',5X,'VA',7X,'VB',8X,'S',
28     *7X,'T')
29     240 FORMAT(5X,19,F10.2,F9.2,F9.2,F9.2)
30     260 FORMAT(5X,19,F10.2,3X,'RUNOUT=',F10.2)
31     270 FORMAT(53X,'V=',F9.3)
32     334 FORMAT(9X,'AVALANCHE DOES NOT STOP'/9X,'CMPUTATNS TERMINATED')
33     445 FORMAT(9X,'NO. OF SGMTS > IMAX'/9X,'COMPUTATIONS TERMINATED')
34     556 FORMAT(9X,'FLO VELCTY NEGLIGBLE'/9X,'COMPUTATNS TERMINATED')
35     667 FORMAT(9X,'VSCSITY NOT SPECIFD'/9X,'CMPUTATNS TERMINATED')
36     C
37     C INITIAL COMPUTATIONS
38     100 WRITE(6,60) H,SML,XNU,YMU
39     IF(VNU(1).LE.0.0) GO TO 666
40     VA=0.0
41     VMAX=0.0
42     IT=I-1
43     I=1
44     S=0.0
45     T=0.0
46     120 IM=1
47     SUL=SML
48     YUU=YMU
49     IF(I.EQ.1) WRITE(6,210)
50     A=G*SIN(THETA(I))-G*YUU*H*COS(THETA(I))
51     B=VNU(I)/H**3
52     E=EXP(-2.0*B*SEGL(I))
53     P=VA*VA*E+A*(1.0-E)/B
54     IF(VA.EQ.0.0.AND.P.LT.25.0) GO TO 130

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55      IF (VA.NE.0.0.AND.P.LT.25.0) GO TO 150
56      130  VB=SQRT(P)
57      IF (VMAX.LT.VB) VMAX=VB
58      S=S+SEGL(I)
59      T=T+2.0*SEGL(I)/(VA+VB)
60      WRITE(6,240) I,VA,VB,S,T
61      Q=THETA(I)-THETA(I+1)
62      VT=0.05*VMAX
63      IF (VA.LT.VT.AND.VB.LT.VT) GO TO 190
64      VA=VB*COS(Q)
65      IF (Q.LE.0.0) VA=VB
66      I=I+1
67      IF (I.GT.IT) GO TO 333
68      GO TO 120
69      C    SUB-SEGMENT COMPUTATIONS
70      150  IN=INT(SEGL(I)/SUL)
71      V=VA
72      SL=SEGL(I)
73      SEGL(I)=0.0
74      IF (IN.LE.1) GO TO 180
75      170  IF (V.LE.5.0) B=XNLI*VNU(I)/H**3
76      IF (V.GT.5.0) B=VNU(I)/H**3
77      E=EXP(-2.0*B*SUL)
78      P=V*V*E+A*(1.0-E)/B
79      VN=SQRT(P)
80      IF (VMAX.LT.VN) VMAX=VN
81      VT=0.05*VMAX
82      SEGL(I)=SEGL(I)+SUL
83      IF (VN.LT.VT.AND.V.LT.VT) GO TO 190
84      IF (IM.EQ.IN+1) GO TO 130
85      V=VN
86      WRITE(6,270) V
87      IF (IM.EQ.IN) SUL=SL-SEGL(I)
88      IM=IM+1
89      GO TO 170
90      180  SEGL(I)=SL
91      GO TO 130
92      190  S=0.0
93      DO 200 J=1,I
94      200  S=S+SEGL(J)
95      WRITE(6,260) I,VA,S
96      GO TO 555
97      C    ERROR MESSAGES
98      333  WRITE(6,334)
99      GO TO 999
100     444  WRITE(6,445)
101     GO TO 999
102     555  WRITE(6,556)
103     GO TO 999
104     666  WRITE(6,667)
105     999  CONTINUE
106     STOP
107     END

```

Summary

Three computer programs, useful in analysis of snow avalanche runout prediction, have been compared by analysis of The Ironton Park avalanche path. Two of the programs, AVALNCH with fast-stop, and ACCEL have been used previously in typical avalanche analyses. Program BIEQ, and a modified version of AVALNCH, referred to as the biviscous version, have been developed in the course of this work.

Program AVALNCH, with its' two versions, is the most versatile of the codes, since transient fluid processes can be modeled. The other codes, based upon the Voellmy theory of avalanche flow, incorporate fluid equilibrium-flow equations. Although program AVALNCH has greater versatility, it requires orders-of-magnitude more time to run a path analysis, compared to programs BIEQ and ACCEL. Both of the programs BIEQ and ACCEL use the Voellmy equations; however, program BIEQ incorporates parameter definitions that are based upon recent findings on the mechanics of flowing avalanches. In taking account of these mechanics processes in writing program BIEQ, a reduction has been obtained in the variation of parameters in order to model different avalanche cases. As determined from numerical evaluation, the parameter that has strongest influence on avalanche runout is the snow depth. The primary change in program BIEQ is to represent the effects of friction and viscous drag as functions of the snow depth. Friction is made proportional to the depth h , and viscous drag to h^{-3} , based upon physical arguments. Then by selection of one set of values of the proportionality coefficients the Ironton Park avalanche runout for snow depths of 2.0, 1.5, 1.0 and 0.5m was approximated. This is in contrast to the need for different valued coefficients for each of these cases if use is made of previously developed versions of the Voellmy equations. No claim is intended that the depth dependence selected is correct. However, the observed

effect of setting up these relationships in h , in that the variation in parameters in order to match different avalanche runouts is greatly reduced, is encouraging. Pending further checkout of program BIEQ, less sensitivity in parameter selection is expected compared with former versions of analysis methods based upon the Voellmy equations. Obvious improvement of the algorithm would be to incorporate snow depth changes as the avalanche advances along its' path. Program AVALNCH does this, but since it is a 2-dimensional code, lateral expansion or contraction of a flow is not accounted for. The only redeeming aspect of this flow depth variation problem is the tendency in viscous fluid dynamics that as the flow increases in depth friction effects decrease and viscous effects increase, and visa versa as flow depth decreases. Thus, if account is taken of both friction and viscous processes, then they have an interactive balancing effect with changing flow depths. In the case of the Ironton Park avalanche path the width of avalanche runout is nearly constant, which simplifies the modeling problem.

Incorporated in program BIEQ is a version of the physical condition of material locking, which has been observed with snow flow. The representation used for material locking is that viscosity of the flow begins to increase at a flow speed of $\sqrt{f} = 8.0 \text{ms}^{-1}$, and exponentially increases as the speed decreases. The material locking algorithm that is used in BIEQ is selected based upon a single segment equilibrium modeling of snow flow, and is only one of many that currently could be selected. With further check-out of program BIEQ, more rational representation of the material locking algorithm should become evident.

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雪崩走出予知用コンピュータープログラム

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この報告書は雪崩走出の解析用として、新庄支所の計算機システムに使えるように書き換えられた計算機用プログラム三種の機能と応用とをとりまとめたものである。雪の流動現象のモデル化の際の流体力学原理としては、このコンピューターコード三つのうちの二つは平衡流体動力学方程式に基づいたものを利用している。三番目のものは二次元非圧縮性境界層理論に基づいた過渡的粘性効果に準拠している。上述の二つの流体力学に基づいたコードのうち、一つは既に開発されていたものであるが、それには一定範囲の摩擦係数と粘性係数とが含まれており、これらは個々の雪崩の型により大きく異なるため経験を積んだ者以外にはその利用は仲々困難であった。二番目のコードは新庄支所で開発されたものであり、それは流れの厚さ依存性を含むと同時にまたその流れる物質の性質を表わす諸係数を明確化する際に、閉塞の性質をも含んでいるものである。それぞれのコードのリスト、すなわち計算機へのデータインプット操作のフォーマットと順序がこの報告書に含まれている。あらかじめ予想された速度プロファイルと走出距離とそれぞれの計算コードを用いて求められたものと比較が、一つの雪崩走路（アメリカ・コロラド州のアイロントン公園）に対してなされた。速度プロファイルはそれぞれのコンピューターコードで異なりはしたものの、走出距離については適当なパラメーターを選ぶことによりうまく適合させることができた。

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